EMISSION OF SHORT-LIVED HALOCARBONS BY SELECTED TROPICAL MARINE PHYTOPLANKTON

LIM YONG KIAN

INSTITUTE OF GRADUATE STUDIES UNIVERSITY OF MALAYA KUALA LUMPUR

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LIM YONG KIAN

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ABSTRACT

Biogenic volatile halocarbons contribute free halogen radicals to the troposphere and stratosphere, and may play a role in the catalytic destruction of the ozone layer. The contributions and significant impacts of biogenic halocarbon emissions from the tropics are relatively unknown and are of particular interest due to the prevalence of strong convective forces at the tropics and climate change. Of the marine biogenic sources, the marine microalgae (phytoplankton) inhabiting the oceans that cover 70% of the Earth's surface, make them a significant source of the short-lived halocarbons. A change in the environment may affect the emission of halocarbons. In this study, the effects of lifestage and irradiance were investigated. Using controlled laboratory experiments, three selected tropical marine phytoplankton were investigated for emission of halocarbons. Three phytoplankton species were grown in flask cultures and sampled for halocarbon emissions during different growth stages of the batch cultures. Growth was estimated using chlorophyll-a and cell number. Halocarbons were measured using a two-syringe collection system followed detection using a GC-MS equipped with a purge and trap system. The phytoplankton were found to emit a suite of short-lived halocarbons, namely CHBr₃, CH₃I, CHCl₃, CHBr₂Cl and CH₂Br₂ at different growth phases. Amphora sp. UMACC 370 was shown to be a stronger halocarbon emitter, especially CH₃I (10.55 – 64.18 pmol mg⁻¹ day⁻¹), than the other two taxa, *Synechococcus* sp. UMACC 371 and *Parachlorella* sp. UMACC 245 $(1.04 - 3.86 \text{ pmol mg}^{-1} \text{ day}^{-1} \text{ and } 0 - 2.16 \text{ pmol mg}^{-1} \text{ day}^{-1}$ ¹, respectively). CH₃I has significantly (p < 0.05) higher emission rate compared to the other detected compounds. Results show that the emissions of detected short-lived halocarbons are species- and growth phase-dependent, highlighting the importance of considering cell physiological conditions when determining gas emission rates. Chlorophyll-a and cell density normalized to emission rate of all five compounds were found to be highly correlated (p < 0.01). The cultures were also exposed to a range of irradiance, 0, 40 and 120 µmol photons m⁻² s⁻¹. The photosynthetic performance (F_v/F_m , maximum quantum yield) of the cultures when exposed to the range of irradiance was used as an indicator of algal cell stress from photosynthesis. F_v/F_m was measured using the Water Pulsed Amplitude Modulated Fluorometer (PAM). Exposure to 120 µmol photons m⁻² s⁻¹ for 12 hours produced significant (p<0.05) decrease in F_v/F_m and increase in halocarbon emissions, especially the release of CH₃I by *Amphora* sp. UMACC 370. The net changes of F_v/F_m , however, were weakly correlated to the significant (p<0.05) changes in overall emission of the five compounds, suggesting that halocarbon emission triggered from oxidative cell stress at higher irradiance may not be directly linked to photosynthesis but instead to mitochondrion respiration, nutrient limitation or a change in lipid composition within the cell membrane.

ABSTRAK

Halokarbon biogenik meruap menyumbang halogen radikal secara bebas ke troposfera dan stratosfera dan mungkin memainkan peranan dalam pemusnahan lapisan ozon. Sumbangan dan kesan nyata pelepasan halokarbon biogenik dari kawasan tropika masih tidak diketahui dan amat diminati disebabkan kelaziman perolakan yang kuat di kawasan tropika dan perubahan iklim. Satu sumber biogenik marin, fitoplankton yang mendiami lautan meliputi 70% daripada permukaan bumi membuat mereka satu sumber halokarbon hayat-pendek yang penting. Perubahan dalam persekitaran tentu boleh mempengaruhi pelepasan halokarbon. Dalam kajian ini, kesan peringkat pertumbuhan dan sinaran cahaya telah disiasat. Menggunakan eksperimen makmal terkawal, tiga fitoplankton marin tropika yang terpilih telah disiasat untuk mengaji pelepasan halokarbon. Tiga spesies fitoplankton telah ditumbuh dalam kelalang dan disampel untuk pelepasan halokarbon semasa peringkat pertumbuhan yang berbeza. Pertumbuhan dianggar menggunakan klorofil-a dan bilangan sel. Halokarbon diukur menggunakan sistem pengumpulan dua picagari dan diikuti pengesanan halokarbon menggunakan GC-MS dilengkapi dengan sistem pembersihan dan perangkap. Fitoplankton didapati melepaskan satu set halokarbon hayat-pendek, iaitu CHBr3, CH3I, CHCl3, CHBr2Cl dan CH2Br2 di fasa pertumbuhan yang berbeza. Amphora sp. UMACC 370 telah ditunjukkan sebagai halocarbon pemancar yang lebih kuat, terutamanya CH₃I (10.55 – 64.18 pmol mg⁻¹ day⁻ ¹), berbanding dengan dua taksa yang lain, iaitu Synechococcus sp. UMACC 371 dan *Parachlorella* sp. UMACC 245 $(1.04 - 3.86 \text{ pmol mg}^{-1} \text{ day}^{-1} \text{ and } 0 - 2.16 \text{ pmol mg}^{-1}$ ¹ day⁻¹, masing-masing). CH₃I mempunyai signifikan (p < 0.05) kadar pelepasan yang lebih tinggi berbanding dengan sebatian dikesan lain. Keputusan menunjukkan bahawa pelepasan halokarbon hayat-pendek yang dikesan adalah spesies- dan pertumbuhan fasa pergantungan, menonjolkan kepentingan untuk mempertimbangakan sel keadaan fisiologi apabila menentukan kadar pelepasan gas. Klorofil-a dan ketumpatan sel

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dinormalkan kadar pelepasan kesemua lima kompaun telah didapati berkait-rapat (p<0.01). Fitoplankton juga didedahkan dengan pelbagai sinaran, 0, 40 dan 120 µmol foton m⁻² s⁻¹. Prestasi fotosintesis (F_v/F_m, hasil kuantum maksimum) daripada tumbuhan yang terdedah kepada julat sinaran digunakan sebagai penunjuk tekanan sel-sel. Pulsed Amplitude Modulated Fluorometer (PAM) digunakan untuk mengukur F_v/F_m. Pendedahan pada 120 µmol foton m⁻² s⁻¹ selama 12 jam menghasilkan penurunan F_v/F_m. Pendedahan 120 µmol foton m⁻² s⁻¹ selama 12 jam menghasilkan penurunan F_v/F_m dan peningkatan pelepasan halokarbon yang signifikan (p<0.05), terutamanya pembebasan CH₃I oleh *Amphora* sp. UMACC 370. Walaubagaimanapun, perubahan F_v/F_m lemah dikait-rapat dengan perubahan signifikan (p<0.05) pelepasan lima sebatian, menunjukkan bahawa pelepasan halokarbon dicetuskan daripada tekanan oksidatif sel daripada sinaran yang lebih tinggi tidak semestinya diakibatkan fotosintesis tetapi berkemungkinan lebih berkait-rapat dengan pernafasan mitokondrion, had nutrien dan perubahan komposisi lipid dalam membran sel.

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LIST OF SYMBOLS AND ABBREVIATIONS

%	:	Percentage
<	:	Less than
>	:	More than
°C	:	Degree Celcius
ANOVA	:	Analysis of Variance
Ca	:	Chlorophyll-a
CADOI		Algal Collection at The University of North Carolina at
CAROL	•	Chapel Hill
CCMP	:	Culture Collection Marine Phytoplankton
CH ₃ *	:	Radical methane ion
CFC	:	Chlorofluorocarbon
Chl-a	:	Chlorophyll-a
ClONO ₂	:	Chlorine Nitrate
cell mL ⁻¹	:	cell per millimeter
DW	:	Dry Weight
ENSO	:	El Nino-Southern Oscillation
ETR _{max}	:	Maximum Electron Transport Rate
Fo	:	Minimum fluorescence
Fm	:	Maximum fluorescence
F_v/F_m	:	Maximum Quantum Yield
FESEM	:	Field Emission Scanning Microscope
GC	:	Gas Chromatography
GCMS	:	Gas Chromatography-Mass Spectrometer
GHG	:	Greenhouse Gas

Gg Br yr ⁻¹	:	Gigagram Bromine per year
Gg I yr ⁻¹	:	Gigagram Iodine per year
g L ⁻¹	:	gram per Litre
H_2O_2	:	Hydrogen Peroxide
HCl	:	Hydrochloric Acid
I*	:	Radical iodine
ΙΟ	:	Iodine Oxide
μ	:	Micro
µg Chl-a ⁻¹ h ⁻¹	:	Microgram per Chlorophyll a per hour
μ mol photon m ⁻² s ⁻¹	:	Micromole photon per square metre per second
mmol Br y ⁻¹	:	millimole Bromine per year
Mmol yr ⁻¹	:	Megamole per year
mg m ⁻³	:	milligram per volume
m/z	:	Quantifying Ion
MSD	:	Mass Spectrometry Detector
ng min ⁻¹ m ⁻²	:	nanogram per minute per square metre
nm	:	nanometre
NOx	:	Nitrogen Oxide
NPQ	:	Non-Photochemical Quenching
NS	:	Non-Significant
O ₂	:	Oxygen molecule
O ₃	:	Ozone molecule
OBIS	:	Ocean Biogeographic Information System
OFN	:	Oxygen Free Nitrogen
р	:	Pearson Correlation Coefficient
PAM	:	Pulse Amplitude Modulation

pМ	:	picoMolar
pmol g DW ⁻¹ h ⁻¹	:	picomole gram per Dry Weight per hour
pmol L ⁻¹	:	picomole per Litre
pmol mol ⁻¹	:	picomole per mole
pptv	:	part per trillion volume
PSC	:	Polar Stratospheric Cloud
PSII	:	Photosystem II
Prov50	:	Provasali 50 culture medium
r	:	Regression
r^2	:	Coefficient of Detemination
rpm	:	round per minute
SAM	:	S- adenosyl-L-methionine
S.D.	:	Standard Deviation
Sig.	:	Significance
sp.	:	species
SPSS	:	Statistical Package for the Social Sciences
Std.	:	Standard
UMACC	:	University of Malaya Algae Culture Collection
LITEV		The Culture Collection of Algae at The University of Texas
UTEA	•	at Austin
uv	:	Ultraviolet Radiation
Va	:	Volume of acetone
Vc	:	Volume of culture
VOC	:	Volatile Organic Compound
X	:	Halide

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CHAPTER 1: INTRODUCTION

1.1 Background

The group of halogenated compounds known as halocarbons, have received less attention for their contributions to climate change than other greenhouse gases such as carbon dioxide, methane and nitrous oxide. Once in the atmosphere, halocarbons give rise to bromine, chlorine and iodine radicals that can cause the catalytic destruction of ozone, resulting in increased penetration of harmful UV radiation to the Earth's surface (Frederick, 2015). Most of the long-lived halogenated compounds are derived from manmade (anthropogenic) chemicals and are notoriously responsible for the drastic loss of stratospheric ozone. Apart from the contribution of anthropogenic activities, scientists are looking further into natural activities that might result in the release of volatile organic compounds, in order to minimize the uncertainties in the estimation of global halocarbon budget. Marine biogenic sources such as phytoplankton (microalgae) are one of the top contributors of the shorter-lived compounds to the atmosphere as they are widely distributed throughout the euphotic zone of all of the Earth's aquatic environments (Moore, 2003). Recent successes in using algae as feedstocks for biofuel, industrial biomaterials and biopharmaceuticals that have initiated large-scale mass cultivation of the phytoplankton may enhance the release of halocarbons quantitatively. A number of short-lived halocarbon compounds released from oceanic sources such as from the marine phytoplankton of polar and temperate regions as well as seaweeds (macroalgae) had been reported following the discovery of increased levels of iodomethane, CH₃I over kelp beds (Keng et al., 2013). Recent studies suggested the possibility of these biogenic short-lived halocarbons contributing and adding to the stratospheric halogen load.

Though literature on halocarbon emissions by the polar and temperate phytoplankton is available, reports from the tropics are relatively unknown. The transport of short-lived halocarbons to the tropical tropopause is very rapid. This fast ascent of the halocarbons is linked to the occurrence of deep convection that is prevalent in the tropics due to a combination of high insolation and high humidity. This in turn may be modulated by the incidence of El Niño-Southern Oscillation (ENSO) events such as through effects on monsoon dynamics. (Bergman *et al.*, 2012; Navarro *et al.*, 2015; Hossaini *et al.*, 2015). In their analysis of global warming trends, Mora *et al.* (2013) reported that in the next 10 years, and before big temperature, ice-melting shifts are seen in the Arctic, the tropics will suffer "unprecedented" climate change effects.

1.2 Problem Statement

The negative effects of climate change events such as global warming affect the emission (rate) of volatile organohalogen (halocarbons) by marine microalgae. The halocarbons, in turn, can increase the earth's temperature through depletion of the ozone. Hence, it is crucial that some work in the tropical region be done to better understand the local atmospheric chemistry and its contribution to the global scenario.

1.3 Research Questions

- (i) What are the main halogenated compounds emitted by tropical marine microalgae?
- (ii) How do different physiological growth phases of the tropical marine microalgae affect the emission of halocarbons?
- (iii) How does irradiance affect the emission of halocarbons?

1.4 Objectives

The overall objective of this research project is to investigate and understand shortlived halocarbon emission by tropical marine microalgae. The sub-objectives undertaking this research are:

(1) To identify the main halocarbons emitted by tropical marine microalgae

- (2) To profile the halocarbons during the growth cycle of selected microalgal cultures
- (3) To study the effect of irradiance and photosynthetic performance on halocarbon emission by selected tropical marine microalgae in the laboratory

1.5 Thesis outline

This thesis is divided into six chapters described briefly as follows:

- (i) Chapter 1 introduces the background of this study and addresses the associated problems and research questions with appropriate objectives.
- (ii) Chapter 2 defines relevant terms, describe research scopes and previous works by referring to the literature.
- (iii) Chapter 3 provides the experimental design and methodology from the beginning of microalgal culturing to finding out the emission of halocarbons.
- (iv) Chapter 4 presents the experimental results in proper format and validates the model of study.
- (v) Chapter 5 discusses the results in comparison with relevant literature accordingly.
- (vi) Chapter 6 gives an overview of the research presented and concludes the findings and their contribution.

CHAPTER 2: LITERATURE REVIEW

2.1 Halocarbons

2.1.1 Halocarbon chemistry

Halocarbons, also commonly known as halogenated compounds or organohalogens, are molecules that comprise of carbon atoms covalently bonded to one or more halogens such as chlorine (Cl), bromine (Br), iodine (I) and Fluorine (F) in the presence or absence of hydrogen (McMurray, 2008). Halocarbons can be toxic and volatile, and are generally unreactive (Sneader, 2005). Bromo- and iodo-carbons are considered more reactive than chlorocarbons and even more so than fluorocarbons, due to the higher stability of the corresponding halide anions and the stronger single-bonded strength of the latter (McMurray, 2008). Halocarbons can be classified into long-lived and short-lived.

2.1.2 Long-lived halocarbons

Long-lived halocarbons are halogenated compounds with atmospheric lifetimes longer than six months (WMO, 2014). They contribute to the halogen load in the inner atmosphere layer, the stratosphere. This has to do on the account of the fact that longlived halocarbons are not significantly degraded in the troposphere during their transport to the stratosphere since they exist longer than the time needed to move to the stratosphere (WMO, 2014).

Long-lived anthropogenic halocarbons are involved in affecting the atmospheric chemistry. They include, but not limited to, the chlorofluorocarbons (CFCs), hydrochlorofluorocarbons (HCFCs), chloromethane (CH₃Cl), tetrachloromethane (CCl₄) and many other halons (WMO, 2014). Some CFCs compounds such as CFC-11, CFC-12 and CFC-113 have very long lifetimes ranging from 45 years to 100 years (WMO, 2014). Long-lived anthropogenic compounds are predominantly made up of chlorine atoms and they were found to be responsible for the major stratospheric ozone loss (UNEP, 2010).

Volcanic eruptions also release some of the long-lived halocarbons like CCl₄ and CH₃Cl, but the contribution to the halogen loads, overall, is negligible since the chlorines are easily dissolved in water and washed out of the atmosphere in rain (Jordan, 2003).

2.1.3 Short-lived halocarbons

Short-lived halocarbons, also commonly referred as trace gases, are halogenated compounds with atmospheric lifetimes of six months or less (Laube *et al.*, 2008). These trace gases can also be classified into very short-lived halocarbons, depending on their specific range of atmospheric lifetimes. They play a part in contributing and adding free halogen radicals to the troposphere and stratosphere, even though long-lived halocarbons are more notoriously known as the main long-term culprit of halogen radical contributors (Laube, 2008).

Examples of chlorinated short-lived compounds include trichloromethane (CHCl₃), dichloromethane (CH₂Cl₂), tetrachloroethylene (C₂Cl₄) and many more. All chlorinated compounds can be either from anthropogenic or natural sources. For instance, CHCl₃ can be produced naturally by seaweeds as well as marine microalgae (Colomb *et al.*, 2008; Nightingale *et al.*, 1995). C₂Cl₄, an excellent solvent widely used as spot remover and for dry cleaning, are synthetically produced (WMO, 2014).

Brominated short-lived halocarbons include tribromomethane (CHBr₃), dibromomethane (CH₂Br₂) and bromomethane (CH₃Br). Most of the anthropogenic sources of CH₃Br are from the agricultural use during soil fumigation. The remaining anthropogenic sources are form biomass burning, leaded petroleum, industry and structural fumigation (WMO, 2011). The biological production of CH₃Br by marine phytoplankton (microalgae) was reported (Sæemundsdottir & Matrai, 1998). Terrestrial vegetation, seaweeds and marine microalgae were found to be some of the natural sources of CHBr₃ and CH₂Br₂ (Moore *et al.*, 1996; Carpenter *et al.*, 2000). The oceanic emission of CHBr₃ had been estimated to be 430-1400 Gg Br yr⁻¹, and CH₂Br₂ 57-280 Gg Br yr⁻¹ (WMO, 2014).

Short-lived organoiodines such as iodomethane (CH₃I), diiodomethane (CH₂I₂), chloroiodomethane (CH₂CII), bromoiodomethane (CHBr₂I) and iodoethane (C₂H₃I) were some of the widely studied marine-produced volatile iodinated organohalogens. These compounds have received attention in connection to the chemistry of iodine in the atmosphere (Moore, 2003). About 214 Gg I yr⁻¹ of iodocarbons were produced by microbial activity in rice paddies and by the burning of biological materials (Bell & Hsu, 2002). These volatile iodomethanes are broken up in the atmosphere as part of the global iodine cycle (Bell & Hsu, 2002).

CH₃I is thought to be the most abundant and has been of particular interest by which iodine plays a role in new particle formation in the atmosphere (Carpenter *et al.*, 2000; Manley & Dastoor, 1987). Most of the iodinated short-lived compounds are released through natural processes, especially from the ocean (WMO, 2014). Macroalgae and microalgae are some of the natural sources that contribute to iodine load in the atmosphere. Anthropogenic sources of iodinated compounds such as CH₃I may also be found from the production of biomass burning and fumigation (UNEP, 2010; Mead *et al.*, 2008)

2.1.4 Environmental role of halocarbons

Synthesized halocarbons have been emitted in robust quantities into the environment for the past few decades while halocarbons from the natural sources in the ecosystems have long been emitted for millions of years, contributing to the halogen fluxes in the atmosphere (Gribble, 1998). C₂Cl₄, one of the chlorocarbons, was amongst the first synthesized halogenated compounds discovered by Michael Faraday in the early nineteenth century (Faraday, 1821; Clowes, 2014). It was not until the early twentieth century where many halogenated compounds were created and used for industrial purposes. Ever since the first synthesis of fluorocarbon in the laboratory by Frederic Swarts and later on improved synthesis process of the fluorocarbon by Thomas Midgely, CFCs were broadly used for air-conditioning gases, as aerosol propellants and many more (Sneader, 2005; Thompson, 1932; Swarts, 1908).

Halocarbons, especially the brominated and chlorinated, are known to cause ozone layer depletion. (Forster & Joshi, 2005). A single chlorine atom in the stratosphere can destroy many ozone molecules through a catalytic cycle when ultraviolet radiation is present. Bromine can efficiently destroy up to 40-100 times in destroying ozone molecules than chlorine in the stratosphere (Penkett *et al.*, 1995). In the stratosphere, the halogen radicals produced from halogenated source gases react with ozone molecules in the presence of ultraviolet radiation, ultimately resulting in the loss of ozone layer. These halocarbons may be produced by both natural and anthropogenic processes, such as biogenic pathways at the land and ocean surface of the Earth and industrial releases, respectively (WMO, 2014). The halocarbon released are transported to the stratosphere through vertical transport, e.g. deep convective forces (Aschmann *et al.*, 2009).

Halocarbons are also known to involve in the absorption of infrared radiation from the Earth's surface (WMO, 2014). Halocarbon such as CFCs can exert up to 6000 times of Global Warming Potential (GWP) than other greenhouse gases like carbon dioxide (CO₂), making it much more efficient in absorbing the radiation, and also emit the radiation back to the Earth's surface, resulting higher global temperature on Earth's surface and lower part of the atmosphere (IPCC, 2007).

Iodinated compounds, on the other hand, are not directly involved in the depletion of stratospheric ozone. The number of iodine atoms that reach the stratosphere is greatly reduced due to rapid tropospheric loss (WMO, 2014). However, the release of iodine,

mainly through photolysis in the atmosphere, may be involved in the cyclic catalytic destruction of ozone. These iodine form iodine oxides (IO/OIO) rapidly with ozone, influencing the tropospheric oxidizing capacity (McFiggans *et al.*, 2000) and greenhouse gas processing. This, in turn, would affect the halo-chemistry composition in the stratosphere and create unseen reactions that may enhance the depletion of ozone layer.

2.2 Climate change

2.2.1 Causes and effects of modern climate change

Global warming, refers to the rise in average temperatures of the Earth's surface that is primarily due to the anthropogenic use of fossil fuels, resulting in climate change (Day *et al.*, 2011). Global warming will not only result in higher temperature but also accelerated sea-level rise, changes in rainfall and even in the frequency and intensity of tropical storms (Day *et al.*, 2011). Since the beginning of pre-industrial era, human influences have been the dominant detectable influence on climate change (Houghton *et al.*, 2001). Climate change will interact with other human impacts to produce environmental effects greater than with climate change alone and ultimately leave unwanted impacts on the ecosystem (Day *et al.*, 2011). In other words, living organisms on land and in oceans will inevitably be adversely affected by the climate change will affect the distribution and deposition of the ozone concentration in the atmospheric boundary layer (Kinney, 2008; Jacob & Winner, 2009; Watson *et al.*, 2016).

Anthropogenic halocarbons such as CFCs are not only notorious for its ability to destroy the stratospheric ozone, but also well-known for its contribution to global warming as the dominating and effective greenhouse gases (GHGs) (Mactavish & Buckle, 2013; Ramanathan & Feng, 2009). GHGs are atmospheric gases that have the ability to absorb and emit radiation within the thermal infrared range on Earth, causing

what is known as the "greenhouse effect". Some of the major GHGs include anthropogenic carbon dioxide (CO₂), methane and CFCs from human activities, and water vapors from the nature (Kiehl & Trenberth, 1997). Therefore, these anthropogenic halocarbons contribute more to temperature rise on Earth than the CO₂ (Velders *et al.*, 2007).

The sun radiates a net of 240 Watts/m² of energy in the form of ultraviolet (UV) (Scheff & Frierson, 2014), visible and near Infrared Range (IR) to the surface of the Earth after passing through the atmosphere and about half of the solar radiation is absorbed by the Earth's surface, warming up the Earth. Some of the solar radiation is reflected back to the atmosphere at 103 Watts/m² (Godish *et al.*, 2015). Some of the infrared radiation would pass through the atmosphere and then out into space at 240 Watt/m² while the rest of the IR is converted into hear energy and is absorbed and re-emitted back onto the Earth by GHGs (Godish *et al.*, 2015). Therefore, when there is an accumulation of GHGs molecules in the atmosphere, more heat will be trapped and thus warming up the Earth. In other words, the accumulating abundance of long-lived anthropogenic halocarbon like CFCs in the atmosphere does contribute to the global warming.

The emission of anthropogenic long-lived halocarbons is rampant. Measures and precautions were taken by many international communities to reduce the adverse effects of ozone depletion and global warming contributed by associated halocarbons. The progress to ozone recovery is slowly gaining its momentum towards achieving optimal balance of atmospheric chemistry on Earth because the issues of halocarbon emission by anthropogenic sources were addressed and tackled.

Nonetheless, the sources from natural environments, which have been significantly contributing and adding to the existing halogen load in the stratospheric atmosphere, also play a big role in depleting the ozone layer. Because natural sources of halocarbon emitted

from biomass burning and volcanic activities were insignificant and discounted (Deshler *et al.*, 1996; WMO, 2014), scientists had switched their attention to other main halocarbon contributors, such as seaweeds and marine phytoplankton.

2.2.2 Ozone

2.2.2.1 Importance of ozone

The ozone layer is the earth's primary shield against the harmful ultraviolet radiation (UNEP, 2010). The ozone layer protects all life on Earth by absorbing 97% to 99% of the solar ultraviolet radiation (hv), which, if not would undoubtedly damage exposed life forms on Earth's surface. This essentially leads to undesirable conditions such as skin cancer and weakened immune systems; disrupts marine food web; and reduces crop yield (Nash & Newman, 2011).

2.2.2.2 Ozone production

The earth's atmosphere is categorized into several layers. The lowest level layer, the troposphere, extends from the Earth's surface up to about ten kilometers in altitude. The following layer, the stratosphere, continues upwards from ten kilometers to about fifty kilometers connecting to the mesosphere (Fahey & Hegglin, 2011). The ozone layer, discovered in 1913 by a French physicists Charles Fabry and Henri Buisson (Sivasakthivel *et al.*, 2011), is made of up to 90% ozone molecules (O₃) concentrating in the stratosphere (UNEP, 2010).

Stratospheric ozone is formed naturally by chemical reactions involving oxygen molecules (O_2) and sunlight. Solar ultraviolet radiation breaks apart one O_2 to produce two oxygen atoms (2 O). Each of these highly reactive atoms combines with an O_2 to produce tri-oxygen molecule (O_3), that is, the ozone (Fahey & Hegglin, 2011).

$$O_{2} - - - - (hv)_{---->} O + O$$

 $O + O_2 - O_3$

2.2.3 Effects of halocarbons on atmospheric chemistry

Stratospheric ozone depletion was raised in 1971 for the first time, with the concern that supersonic transport aircraft emission of nitrous oxide and water vapor would adversely affect the ozone levels (Poppoff *et al.*, 1978). Later in 1974, Mario Molina, along with his professor, F. Sherwood Rowland, developed the CFCs ozone depletion theory and discovered that the chlorine atoms, produced by the decomposition of CFCs, can catalytically destroy ozone (Molina & Rowland, 1974). It was concluded that CFCs would not break down on Earth's surface or in the troposphere. Instead, CFCs would rise into the stratosphere and remain for several years. From there, intense *uv* radiation would break their bonds, releasing highly reactive chlorine atoms that quickly and repeatedly react with ozone (UNEP, 2010).

- (1) CCl_3F ------ (hv)-----> CCl_2F + Cl (cfc-11)
- (2) $Cl + O_3 ClO + O_2$
- (3) $ClO + O ----> Cl + O_2$

The net reaction, $O + O_3 - O_2 + O_2$

Each chlorine can destroy as many as up to 100,000 molecules before it become inactivated and returned to the troposphere in the form of HCl (Moore & Stanitski, 2014).

In the late 1980s, a massive "hole" in the ozone over Antarctic was discovered. Satellite data showed that the hole, in terms of percentage of O_3 depletion, had been deepening and enlarging every year since 1977 (Hill, 2010). There has been an extreme depletion since 1987 where the area of the hole has widened to the point where it is larger than the Antarctic continent. The hole tended to worsen progressively in terms of how long it lasts to the Antarctic spring (Sparling, 2001).

The culprit of such ozone hole over the Antarctic lies in the conditions of Antarctic winter and spring that are conducive to O_3 destruction. During winter, the Antarctic stratosphere is denitrified (Toon & Turco, 1991); essentially NOx compounds and water vapors present in the dry air are frozen due to the extreme cold and isolation created from polar vortex, forming what is referred as the "polar stratospheric clouds (PSC) (Chipperfield, 2015). The frozen compounds, which are not present in gaseous form, are therefore not available to react with and tie up chlorine. On top of that, reactions that free Cl from relatively stable reservoirs (HCl or ClONO₂) take place faster on surfaces, as provided by PSC's, than they do in gaseous environment (Chipperfield, 2015).

$$HCl + ClONO_2 ---> Cl_2 + HNO_3$$

When Spring starts to kick in with the first return of direct sunlight, the freed form of Cl, as well as the Cl₂, are photolyzed and photo-associated into atomic Cl, giving them the freedom to react rapidly and repeatedly with O₃. The NOx compounds still remain frozen that hence cannot act to form reservoirs such as ClONO₂ thanks to the extreme cold temperature in the early of Spring due to the vortex (NASA, 2009).

Other organohalogens like Bromine (Br), also take part in similar reactions like the chlorine. In fact, bromine atoms, even at lower concentrations, are 50 times more efficient than Cl at attacking O_3 in the chlorine-rich stratosphere (Berg *et al.*, 1983; Penkitt *et al.*, 1995). Br may be responsible for the 20% Antarctic ozone depletion, with 5-10% of the total depletion due to this particular halocarbon, bromomethane (CH₃Br), alone.

The Arctic also experiences ozone depletions, but not as much as those over the Antarctic; losses of ozone over the Arctic has been around 5-10% range while with 50-66% range over Antarctic (Solomon, 1999). This is mainly due to the fact that: 1) the polar vortex over the Arctic during winter is not as intense as that over the Antarctic, 2) there is a shorter time for the critical overlap between cold and first direct sunlight as the Arctic stratosphere warms faster in Spring than that over the Antarctic, and 3) the Arctic doesn't denitrify as completely as the Antarctic stratosphere does (Mohanakuma, 2008; NOAA, 2010).

The issue of O_3 destruction from natural resources like volcanic eruption was raised back in the 1950s, but enough evidence showed that the losses of O_3 and volcanic activities are not correlated (Deshler *et al.*, 1996). Much of the HCl produced by the volcanoes does not make it to the stratosphere and is quickly washed out through the major Cl removal mechanism from the stratosphere (Deshler, 1996). In fact, there was a significant loss of O_3 back in the 1980s but there was no major volcanic activities (WMO, 2007).

Destruction of ozone in the stratosphere also happens via the cyclic chemical reaction (WMO, 2014). The cyclic chemical reaction involves natural occurring species like halogen radicals, nitrogen oxide radicals and hydrogen radicals. Small changes in radical concentrations will cause serious implications on the O₃ as they get regenerated through the ozone-destructing catalytic cycles (Fahey & Hegglin, 2011).

2.3 Introduction to marine microalgae

2.3.1 What are microalgae?

Phytoplankton comprise the microalgae are microscopic plant-like unicellular organisms capable of efficient photosynthesis and biomass production (Tebbani *et al.*, 2014). They comprise a diverse group of prokaryotic and eukaryotic photosynthetic

microorganisms that grow in both freshwater and marine habitats as well as on soil or as epiphytes. (Li *et al.*, 2008). In a balanced ecosystem, microalgae, serving as the primary producers, play an essential role in marine food chains. They serve as food for a wide range of marine species (Helbling & Villafane, 2001).

2.3.2 Distribution and abundance of microalgae

It has been estimated that about 200,000 to 800,000 of microalgae exist; 5000 out of all these are known species of marine microalgae (Hallegraeff, 2003). The distribution and abundance of microalgae species are controlled by abiotic as well as biotic factors in both space and time. Changes in the microbial community can often be difficult to quantify against a background of high temporal and spatial variability. Nonetheless, evidence indicates that increased precipitation and glacial melt from warmer surface ocean temperatures due to global warming reportedly favors dominance of cryptophytes over diatoms in Antarctic coastal waters (Moline & Prézelin 1996; Moline *et al.*, 2004).

Three most important classes of microalgae in terms of abundance include green algae (Chlorophyceae), the diatoms (Bacillariophyceae) and the golden algae (Chrysophyceae) (Carlsson et al., 2007). Cyanophyceae, a special class of microalgae, which is often called blue-green algae, or Cyanobacteria, a phylum of bacteria capable of obtaining energy from sunlight. The blue-green algae are often referred to as Cyanobacteria because they have cell structure and composition similar to those of prokaryotic cells in that they lack cell nucleus and distinctive organelles of eukaryotes, and their structure and chemical composition of the cell wall are the same as those of gram-negative bacteria (Pisciotto *et al.*, 2010). On the other hand, they also possess pigments like of those in eukaryotic algae to carry out photosynthesis (Pisciotto *et al.*, 2010).
Microalgae are widely populated in many different aquatic environments, from freshwater to brackish and marine waters. (Arrigo *et al.*, 2012; Fisherman *et al.*, 2010; Boonyapiwat, 1997).

The north-south trend of decreased calcification of Emiliania huxleyi in the Southern Ocean over the past two decades since 1983/1984 indicates that Emiliania huxleyi populations are migrating polewards (Cubillos et al., 2007) while the red tide dinoflagellate, Noctiluca scintillans are migrating southwards towards the Southern Ocean from Tasmania brought on by a warm-core eddy circulation (McLeod et al., 2012). Species composition and abundance in Albatross Bay, Gulf of Carpentaria, northern Australia examined from 1986-1992 reflects a stable tropical microalgae community in waters without pulses of physical and chemical disturbances (Burfold et al., 1995) as there was no distinct species succession of diatoms. The diatoms were the dominating species at the inshore sites. The proportion of green flagellates increased at the offshore sites and the cyanobacterium genus Trichodesmium and the diatom genera Chaetoceros, Rhizosolenia. Bacteriastrum and Thalassionema dominated the phytoplankton (Burfold et al., 1995). Diatoms contribute around 20% of global primary productivity (Malviya et al., 2016) and are predominantly distributed on the Northern Hemisphere (Hasle & Syvertsen, 1996; OBIS, 2015). However, Malviya et al. (2016) has shown that most diatom genera were seen in all oceanic provinces although their ribotype abundance patterns based on a molecular rarefaction analysis were highly variable. Chaetoceros (both subgenera), Corethron and Fragilariopsis were highly abundant in the Southern Ocean. Attheya, Planktoniella, and Haslea were seen primarily in the South Pacific Ocean and Leptocylindrus was found to be highly abundant in the Mediterranean Sea (Malviya et al., 2016). Based on a significant positive relationship of chlorophyll-a and fucoxanthin pigments, diatoms are found more dominant in terms of numerical abundance than prymnesiophytes in the central eastern Arabian Sea (Roy, 2010). Strzepek & Harrison (2004) suggested that diatoms, and most probably other eukaryotic algal taxa, might have adapted the ability to survive in different underwater light climate between oceanic and coastal waters, enabling them to decrease their iron requirements without compromising photosynthetic capacity. This facilitates the colonization of the open oceans by diatoms.

2.4 Marine biogenic sources of halocarbons

2.4.1 Halocarbon emissions by marine microalgae

The oceans have been recorded to be the main contributor of volatile organohalogens to the atmosphere, but the sources of organohalogens had been unknown except for macrophytic algae (seaweeds), which primarily are confined to the coastal zone (Moore, 2003).

Krysell (1991) reported that pelagic marine algae are a source of bromoform in the surface waters of the Arctic Ocean. Sturges *et al.* (1993) reported that Arctic ice microalgae emit significant quantities of bromoform that may be converted photochemically into active bromine forms. The active form of bromines, in return, is thought to be one of the main causes of the destruction of surface ozone in the Arctic environment during the spring. The estimates of the total annual bromoform release revealed that polar ice algae might actually contribute globally significant amount of organic bromine compounds, comparable with anthropogenic and macrophyte sources (Sturges *et al.*, 1993). This study was followed by investigations of halocarbon emissions by microalgae originating from different climatic zones from the poles to the tropics. A summary of all the studies on halocarbon emissions by marine microalgae isolated from polar, temperate and tropical zones is reported in Table 2.1 below.

	Climate zones	Type of halocarbons emitted																					
	(incubation temperature)		"	4	сI СI	L2	Ξ		$3r_{2}$,	E			-	12			Ļ	Cl	Ū	
Таха		ΉJ	HCL	HBr	HBr	H	CH.C	H،J،	3H4	HaB	TH ³ C	CH,C	2.HsC	SHC	$^{3}Cl_{4}$, HsC	ZAHAC	λΗς	3H-I	HJB	HBr	HJB	Reference
Smeeten multiple 271	Transian (25%C)	9	9	U	9	9	9	<u> </u>	9	9	9	4	9	4	9	9	9	9	<u> </u>	9	<u> </u>	<u> </u>	Line of all (consulting all)
Synechococcus sp. UMACC 3/1	Tropical (25°C)																						Lim <i>et al</i> . (unpublished)
Chlorella sp. UMACC 245	Tropical (25°C)																						Lim et al. (unpublished)
Amphora sp. UMACC 370	Tropical (25°C)																						Lim et al. (unpublished)
Mediopyxis helyxis	Polar/ Temperate																						Thorenz et al. (2014)
Porosira glacialis	Polar/ Temperate								/														Thorenz et al. (2014)
Thalassiosira sp.	Polar (2- 4 °C)																						Hughes et al. (2013)
Prochlorococcus marinus CCMP 2389	Temperate (22 °C)																						Hughes et al. (2011)
Synechococcus sp. CCMP 2370	Temperate (22 °C)		0																				Hughes et al. (2011)
Prochlorococcus marinus CCMP 1986	Temperate (20-21°C)																						Brownell et al. (2010)
Synechococcus sp.CCMP 2370	Temperate (20-21°C)																						Brownell et al. (2010)
Calcidiscus leptoporus AC365	Sub-tropical (20- 25 °C)																						Colomb et al. (2008)
Emiliania huxleyi CCMP 371	Sub-tropical (20- 25 °C)																						Colomb et al. (2008)
Phaeodactylum tricornutum	Sub-tropical (20- 25 °C)																						Colomb et al. (2008)
Chaetoceros neogracilis CCMP 1318	Sub-tropical (20- 25 °C)																						Colomb et al. (2008)
Dunaliella tertiolecta	Sub-tropical (20- 25 °C)																						Colomb et al. (2008)
Emiliania huxleyi CCMP 379	Temperate (15 °C)																						Hughes et al. (2006)

Table 2.1: Types of halocarbon emitted by cultures of marine phytoplankton isolated from different climatic zones.

Tetraselmis sp. CCMP 961	Temperate (15 °C)											~	Hughes <i>et al.</i> (2006)
Thalassiosira pseudonana CCMP 1335	Temperate (15 °C)												Hughes et al. (2006)
Porphyridium purpureum CCAP 1380/3	Temperate (22 °C)												Scarratt & Moore (1999)
Guillardia theta CCMP 327	Temperate (22 °C)									Л			Sæmundsdottir & Matrai (1998)
Hemiselmis rufescens CCMP 439	Temperate (22 °C)												Sæmundsdottir & Matrai (1998)
Chaetoceros diversum A1299	Temperate (22 °C)							C					Sæmundsdottir & Matrai (1998)
Chaetoceros atlanticus CCMP 161	Polar (4°C)												Sæmundsdottir & Matrai (1998)
Amphidinium carterae CCMP 1314	Temperate (22 °C)												Sæmundsdottir & Matrai (1998)
Crypthecodinium cohnii CCMP 316	Temperate (22 °C)	\square											Sæmundsdottir & Matrai (1998)
Prorocetrum micans CCMP 1589	Temperate (22 °C)												Sæmundsdottir & Matrai (1998)
Pycnococcus provasolii CCMP 1203	Temperate (22 °C)												Sæmundsdottir & Matrai (1998)
Pavlova sp. CCMP 617	Temperate (22 °C)												Sæmundsdottir & Matrai (1998)
Phaeocystis sp. CCMP 628	Temperate (22 °C)			Γ									Sæmundsdottir & Matrai (1998)
Pavlova gyrans CCMP 608	Temperate (15 °C)												Sæmundsdottir & Matrai (1998)
Pavlova lutheri CCMP 1325	Temperate (15 °C)												Sæmundsdottir & Matrai (1998)
Pleurochrysis carterae CCMP 645	Temperate (15 °C)												Sæmundsdottir & Matrai (1998)
Synechococcus bacillaris CCMP 1333	Temperate (22 °C)												Sæmundsdottir & Matrai (1998)
Thalassiosira pseudonana CCMP 1015	Temperate (15 °C)												Sæmundsdottir & Matrai (1998)
Chaetoceros sp. CCMP 208	Polar (4°C)												Sæmundsdottir & Matrai (1998)
Synedra minuscula CCMP 845	Polar (4°C)												Sæmundsdottir & Matrai (1998)
Tetraselmis levis CCMP 896	Temperate (15 °C)												Sæmundsdottir & Matrai (1998)

Emiliania huxleyi CCMP 373	Temperate (22 °C)												Sæmundsdottir & Matrai (1998)
Navicula sp. CCMP 547	Temperate (18 °C)										Ŷ	T	Manley & de la Cuesta (1997)
Nitzschia sp. CCMP 580	Polar (5 °C)												Manley & de la Cuesta (1997)
Phaeocystis sp. CCMP 1521	Temperate (18 °C)							T					Manley & de la Cuesta (1997)
Porosira glacialis CCMP 651	Polar (5 °C)												Manley & de la Cuesta (1997)
Thalassiosira pseudonana CCMP 1335	Temperate (18 °C)												Manley & de la Cuesta (1997)
Nitzschia punctata UTEX 2041	Temperate (18 °C)												Manley & de la Cuesta (1997)
Chaetoceros neogracile CCMP 1317	Temperate (18 °C)				ſ								Manley & de la Cuesta (1997)
Skeletonema pseudonana UTEX 2308	Temperate (18 °C)												Manley & de la Cuesta (1997)
Pycnococcus provasolii CCMP 1203	Temperate (18 °C)												Manley & de la Cuesta (1997)
Dunaliella parva UTEX 1983	Temperate (18 °C)												Manley & de la Cuesta (1997)
Coccolithus pelagicus CCMP 299	Temperate (18 °C)												Manley & de la Cuesta (1997)
Emiliania huxleyi CCMP 370	Temperate (18 °C)												Manley & de la Cuesta (1997)
Isochrysis galbana CCMP 1323	Temperate (18 °C)												Manley & de la Cuesta (1997)
Synechococcus sp. CCMP 1334	Temperate (18 °C)												Manley & de la Cuesta (1997)
Gymnodinium sp. CAROL 3290	Temperate (18 °C)	/											Manley & de la Cuesta (1997)
Phaeodactylum tricornutum	Temperate (18 °C)												Scarratt & Moore (1996)
Phaeocystis sp.	Temperate (18 °C)												Scarratt & Moore (1996)
Thalassiosira weissflogii	Temperate (18 °C)												Scarratt & Moore (1996)
Nitzschia sp. CCMP 580	Polar (6 °C)												Moore <i>et al.</i> (1996)
Nitzschia arctica CCMP 1116	Polar (6 °C)												Moore <i>et al.</i> (1996)

Nitzschia seriata	Polar (6 °C)											Moore <i>et al.</i> (1996)
Porosira glacialis CCMP 651	Polar (6 °C)			T								Moore <i>et al.</i> (1996)
Navicula sp. CCMP 545	Polar (6 °C)											Moore <i>et al.</i> (1996)
Navicula sp. CCMP 546	Polar (6 °C)									7		Moore <i>et al.</i> (1996)
Porosira glacialis	Polar (4 °C)			Γ								Tait & Moore (1995)
Nitzschia seriata	Polar (4 °C)											Tait & Moore (1995)
Nitzschia sp. CCMP 580	Polar (4 °C)											Tait & Moore (1995)
Odontella mobiliensis	Temperate (20 °C)											Tait & Moore (1995)
Phaeodactylum tricornutum	Temperate (20 °C)											Tait & Moore (1995)
Thalassiosira weissflogii	Temperate (20 °C)											Tait & Moore (1995)
Nitzschia sp. CCMP 580	Polar											Moore <i>et al.</i> (1994)
Porosira glacialis	Polar		٠									Moore <i>et al.</i> (1994)
Nitzschia seriata	Polar											Moore <i>et al.</i> (1994)
Nitzschia sp. CCMP 580	Polar (6 °C)											Tokarczyk & Moore (1994)
Porosira glacialis CCMP 651	Polar (6 °C)											Tokarczyk & Moore (1994)
Nitzschia pungens	Polar (6 °C)	7										Tokarczyk & Moore (1994)
Nitzschia seriata	Polar (6 °C)			T								Tokarczyk & Moore (1994)
Thalassiosira gravida CCMP 986	Polar (6 °C)											Tokarczyk & Moore (1994)
Odontella mobiliensis CCMP 596	Polar (6 °C)							1				Tokarczyk & Moore (1994)
Chaetoceros sp. CCMP 208	Polar (6 °C)			T								Tokarczyk & Moore (1994)
Isochrysis galbana CCMP 1323	Polar (6 °C)			T								Tokarczyk & Moore (1994)

Nitszchia stellata	Polar (0 -1.5 °C)				Sturges et al. (1992)
Porosira pseudodenticulata	Polar (0 -1.5 °C)			$\nabla \Psi$	Sturges et al. (1992)
	· · ·				
= Detected	=Not-detected = Not	reported			

All studies on halocarbon emission are associated only with the marine microalgae. This is largely due to the presence of dissolved halogens in the ocean seawater or synthesized marine mediums such as the F/2 (Guillard & Ryther, 1962) and Prov 50 (CCMP 1996) that makes it possible for the microalgae to methylate and produce halocarbons. *Porosira glacialis*, a diatom, was reported to emit 1000 pmol L⁻¹, 1.0 nmol L^{-1} and 2250 pM (equivalent to 2250 pmol L^{-1}) of CHBr₃ (Tokarczyk & Moore, 1994; Moore et al., 1995; Moore et al., 1996). Furthermore, there was evidence of CH₃I production by Emiliania huxleyi, a Chrysophyte, which was reported to emit CH₃I at 0.35 pmol L⁻¹ (Hughes et al., 2006) and 1.8 pptv (Colomb et al., 2008) contrary to Manley & de la Cuesta (1997), who did not observe CH₃I emission by E. huxlevi. Scarratt & Moore (1999) reported a 40 times faster CH₃I production by the red microalga Porphyridium purpureum than that found by Manley & de la Cuesta (1997) for Porosira glacialis. Colomb et al. (2008) reported that microalgae also emitted anthropogenic chlorinated organic compounds, namely chloroform (CH₃Cl), dichloromethane (CH₂Cl₂), trichloroethylene (C₂HCl₃), tetrachloroethylene (C₂Cl₄), chlorobenzene (C₆H₅Cl) and dichlorobenzene (C₆H₄Cl₂). *Phaeodactylum tricornutum*, a diatom was reported to emit 1750 pptv and 900 pM of CH₃Cl, by Colomb et al. (2008) and Tait & Moore (1995) respectively.

In the early studies of halocarbon emission directly from open-surface waters, there were difficulties in confirming if the elevated levels of bromoform and dibromomethane in Arctic ice cores from Resolute Bay, were contributed by organisms other than the phytoplankton, namely the zooplankton and bacteria (Moore & Tokarczyk, 1993), . This is also a concern in laboratory studies when non-axenic microalgal cultures were used. Manley & de la Cuesta (1997) reported that the associated bacteria in phytoplankton cultures (*Nitzschia* sp., *Navicula* sp., *P. glacialis* and *Phaeocytosis* sp.) did not contribute to the emission of methyl iodide. If the bacteria were the only or main producers of CH₃I,

there would be a continued increase in the amount of CH_3I during stationary phase of algal growth where only bacterial numbers continued to increase. In addition, if CH_3I was, in fact, produced by the bacteria, the rate of CH_3I production normalized to bacterial cells should have remained constant, and not dropped dramatically. This was confirmed by Tokarczyk & Moore (1994) and Moore *et al.* (1995).

Despite the report of weak CH₃I production by undefined microbial populations obtained from decaying kelp tissue (Manley & Dastoor, 1987), and marine bacteria are capable of producing CH₃I in the ocean (Amachi *et al.*, 2001), there has not been any direct evidence showing that marine bacteria are involved in other halocarbon production in either the ocean seawater or the laboratory.

Granfors *et al.* (2013) investigating the role of young Arctic sea ice in halocarbon cycling, reported that halocarbon levels were increased by microorganisms inhabiting the ice. Heterotrophic bacteria were evenly distributed in the ice and was probably more responsible for contributing the halocarbon than the microalgae that were found in the lower layers of the ice. CH₃I production by bacterial aggregates (Asare *et al.*, 2012) and production of CHBr₃, CH₃I, CH₂Br₂ and CHBr₂Cl by cyanobacteria (Karlsson *et al.*, 2008) had been previously reported. In the latter study, the halocarbons were measured during a cyanobacteria bloom in summer in the Baltic sea; with production rates up to 0.3 pmol [μ g chl a]⁻¹ h⁻¹ at midday. Hughes *et al.* (2013) investigated the role of a diatom *Thalassiosira* isolated from the coastal waters of the western Antarctic Peninsula in CHBr₃ and CH₂Br₂ production. Production of CHBr₃ was observed to be linked with a primary metabolic process, while bacteria contributed to inhibition of the compound.

Abrahamsson *et al.* (2004) monitored halocarbon levels in 24 hr stations in the Atlantic part of the Southern Ocean and reported that bromochloromethane, tribromomethane, trichloroethene and diiodomethane were dominant. In an attempt to correlate the

compounds with pigments, their results showed that the haptophytes were identified to be linked to the iodinated compounds. Halocarbon levels were also found to peak twice in the diurnal cycle; once at midday when photosynthesis was most active and in the evening. This had been reported by Ekdahl *et al.* (1998) to be related to respiration in the dark period. The Arabian Sea is a very productive area in the Indian Ocean and a study on the halocarbon production by the marine microalgae showed that halocarbon production was strongly correlated to pigment (Roy, 2010). CHCl₃, CH₂Br₂, CHBr₃ and CCl₄ were the most abundant with CHCl₃ significantly correlated with fucoxanthin suggesting the importance of diatoms and prymnesiophytes in the Arabian Sea. Seasonality in halocarbon production was observed with higher production during the summer monsoon induced upwelling that resulted in higher microalgal productivity.

Smythe-Wright *et al.* (2006) reported that high CH₃I concentrations of up to 45 pmol L^{-1} near lower latitude (20°N to 60°N) of Atlantic and Indian oceans correlates well with the abundance of dominating species, *Prochlorococcus* sp.. Solely from this marine source, approximately 5.3 x 10¹¹ g I yr⁻¹ of global ocean flux of iodine was reported to contribute to the marine boundary layer, putting up a large fraction of the previously estimated total global flux of iodine (10¹¹-10¹² g I yr⁻¹) (O'Dowd *et al.*, 2002). Nonetheless, a laboratory-based study led by Brownell *et al.* (2010), through extrapolating the production of CH₃I by similar species, *Prochlorococcus Marinus* to a global scale, reported an average global production rate of CH₃I; 0.6 Mmol yr⁻¹. This production rate that accounted for 0.03% of the global marine production, on the contrary, suggests that *Prochlorococcus* sp. is not a significant marine source of contributing CH₃I to the atmosphere.

A seasonal study of tropospheric volatile organic compounds (VOC) at the South Pole and selected Antarctic sites, revealed that while the longer lived species of anthropogenic origins (eg. alkyl nitrate) were more abundant in late winter, the compounds of oceanic origin like the bromoform and methyl iodide, were observed in early winter, after the summer peak in biological activity but before oceanic productivity decreased during winter months (Bayersdorf *et al.*, 2010). Bromoform levels correlated with alkyl nitrate, and may be influenced by similar factors linked to seasonality of marine emissions, sea ice extent; atmospheric removal via OH oxidation; and seasonal differences in transport efficiency of marine air masses.

2.4.2 Mechanisms behind the halocarbon emissions

Two pathways, in terms of monohalomethanes and polyhalomethanes, are involved in the production of halocarbons by phytoplankton.

(i) Monohalocamethane

Monohalogenated compounds are produced by a process of methylation of its corresponding halide ion (Manley, 2002). The direct way of forming monohalomethanes is through the transfer of a methyl group to a halide catalyzed by an enzyme. This mechanism requires a methyl donor, for instance, S- adenosyl-L-methionine (SAM) and a halide ion methyl transferase to transfer the methyl group over to the halide ion. Different types of methyl transferase show different kinetics and substrate specificities (Harper, 2000). Some enzymes are capable of catalyzing several halocarbons simultaneously (Itoh *et al.*, 1997). In general, in decreasing order of reactivity are iodide, bromide and chloride, consistent with their decreasing nucleophilicity (Wuosmaa & Hager, 1990).

Several studies of halomethane production from the methyl-transferase system were reported (Wuosmaa & Hager, 1990; Saini *et al.*, 1995; Hughes *et al.*, 2011). Hughes *et al.*, (2011) proposed that the production of methyl iodide (CH₃I) by *Prochlorococcus marinus* (CCMP 2389) could also be due to the iodination of inorganic iodine species, such as atomic iodine (I*) and the breakdown of higher molecular weight organic iodine (Fenical, 1982) and by radicals (CH₃*) or methyl group arising from photochemical reactions (Moore & Zafiriou, 1994). Hughes *et al.* (2011) found connections between the CH₃I production rate with cell stress, increased cell membrane permeability and cell lysis. It was suggested that the cells' response to limiting conditions enhanced the production of CH₃I or that the loss of membrane integrity increased the release of CH₃I precursors to the medium.

(ii) Polyhalomethanes

The formation of polygenated compounds involves enzymatic halogenation via haloperoxidase (Theiler *et al.*, 1978). The enzymatic biohalogenation process involved two steps. Firstly, the haloperoxidase enzyme catalyzes the oxygen-based oxidant co-substrate to react with X^- (X= Cl, Br or I) to produce electrophilic halogen species like XO⁻or R₂NX. Subsequently, the electrophilic halogen species attacks a carbonyl activated methyl group and substitute a hydrogen atom. The next H-atom substitution then leads to the formation of di- or trihalogenated methanes, and so on (Theiler, 1978; Kline *et al.*, 2000).

Haloperoxidase are characterized by their ability to oxidize halogen anions. For instance, chloroperoxidases can use chloride, bromide and iodide; bromoperoxidases can use bromide and iodide whilst iodoperoxidases can only use iodide (Moore *et al.*, 1996; Urhahn & Ballschmiter, 1998).

A study from Moore *et al.* (1996) reported the presence of bromoperoxidase in *Nitzschia* sp. (CCMP 580), with the ability to produce brominated compounds; CHBr₃, CH₂Br₂ and iodinated compounds; CH₂I₂ and CH₂CII. On the other hand, the presence of iodoperoxidase was reported, revealing the production of only the iodinated compounds.

Marine algae are a rich source of volatile halogenated metabolites, especially halomethanes that make algae contribute substancially to the global budget of such molecules. Nontheless, it is often very challenging to address the physiological and ecological interactions for such metabolites that are ubiquitous in algae and their environment. This is especially true for marine phytoplankton where their overall concentrations in the water column might be relevant, and locally elevated amount of metabolites in the immediate vicinity of the producing cells (Paul & Pohnert, 2011). However, studies have indicated that the release of volatile halocarbons by marine phytoplankton has more to do with defence mechanism from herbivory predators (Bravo-Linares & Mudge, 2009; Paul & Pohnert, 2011). Due to the presence of halogen, the halogenated metabolites often have exceptionally high biological activities, not only can aid in chemical defense but also can act as antifouling agents for the producing algal cells (Pau & Pohnert, 2011).

2.4.3 Significance of tropical emission of biogenic halocarbons

The intense tropical convection in the tropics, especially that over the oceans, have been suggested as being responsible for the rapid transport of compounds into the stratosphere (Schauffer *et al.*, 1999; Levine *et al.*, 2007; Laube *et al.*, 2008; Fueglistaler *et al.*, 2009; Brinckmann *et al.*, 2012; Hossaini *et al.*, 2015). The recent years have seen a growing interest in the monitoring of atmospheric compounds, including the short-lived halocarbons in the western Pacific/Southeast Asian region, characterized by high primary productivity (Sherman & Hempel, 2009). The convection is strongest here especially during the Northern Hemisphere winter, and can lift the surface oceanic emissions to the tropical tropopause layer which refers to the transition layer between the troposphere and stratosphere (Fueglistaler *et al.*, 2009; Robinson *et al.*, 2014). Measurements have been taken via whole-air sampling, short-term deployment of gas chromatograph devices and on-board measurements during research cruises. Robinson *et al.* (2014) was the first to report continuous measurement from instruments based in Sabah, Borneo.

The contribution of the coastal primary producers, the macroalgae or seaweeds, to the CHBr3 and CH2Br2 emissions in the Southeast Asian region, was estimated by Leedham et al. (2013) to a range of 6-224 mmol Br yr⁻¹ that was lower than that (180-350 mmol Br yr⁻¹) reported by Pyle *et al.* (2011a), although the definition of SEA was not similar. In extrapolating further to the global scenario, where tropical oceans are estimated to contribute 75% of global halocarbon budget, the very low values from Leedham et al. (2013) suggest that emissions from the open oceans (phytoplankton) may be more important than the coastal emissions. Mohd Nadzir et al. (2014) using data collected during a research cruise over the Straits of Malacca, the South China Sea and the Sulu-Sulawesi Sea in 2009, estimated a regional emission of 63 Gg yr⁻¹ CHBr₃ for the Southeast Asian region. CHBr₃ was the most abundant bromocarbon, ranging from 5.2 pmol mol⁻¹ in the Straits of Malacca to 0.94 pmol mol⁻¹ over the open ocean. Robinson et al. (2014) reported the biogenic short-lived CHBr₃, CH₂Br₂ and CHI₃, from a 15-month survey from the rainforest (Danum) and coastal (Kunak) sites first reported by Pyle et al. (2011b). CHBr3 and CH2Br2 concentrations did not show seasonal variations, but shortterm variations were evident in the coastal site related to the marine sources including coastal seaweed and both coastal and oceanic phytoplankton. They also concluded that the Southeast Asian region may not be as strong a contributor of brominated short-lived halocarbons as previously reported.

Manley et al. (1992) made a comparison on the production rates of bromoform per unit per biomass of microalgae and macroalgae. They concluded that even though the phytoplankton releases were 10-100 times lower, these unicellular organisms had the potential to be a vital source of volatile halocarbons because they occupy an ocean area of about 200 times larger than that occupied by the macroalgae. Of the halocarbons identified from the algae, chloroethane (C_2H_5Cl) , trichloroethane $(C_2H_3Cl_3)$, dichloroethane $(C_2H_4Cl_2)$ and tetrachloroethane (C_2Cl_4) normally associated with anthropogenic origins, have been reported for five cosmopolitan marine phytoplankton, Calcidiscus leptoporus, Emiliania huxleyi, Phaeodactylum tricornutum, Chaetoceros neogracilis and Dunaliela tertiolecta (Colomb et al., 2008) but not in seaweeds. The two diatoms P. tricornutum and Ch. neogracilis were the strongest emitters of methyl bromide. The latter three species are mass cultivated for aquaculture feed and as a feedstock for biofuel production. Two diatoms, Mediopyxis heylysia and Porosira glacialis also released dibromopropane (C₃H₆Br₂) which has not been detected in seaweeds (Thorenz et al., 2014). Although there has been more research conducted on the halocarbon emissions by seaweeds than phytoplankton, current available literature allows a preliminary estimation of type and range of halocarbons identified. In general, 39 compounds (10 iodinated, 7 brominated, 13 chlorinated and 9 mixed halocarbons) have been reported as being emitted by the clonal-cultured phytoplankton (18) and fieldcollected seaweeds (34), with 13 compounds in common; giving a Sørensen's Coefficient of Similarity of 0.50, suggesting that 50% of the detected halocarbon species present in seaweeds are also present in phytoplankton. As 0.5 in similarity also indicates 0.5 in dissimilarity, this suggests a potential emission of previously detected compounds by seaweeds to be seen in phytoplankton, and vice versa.

2.4.4 Factors affecting halocarbon emissions by microalgae

2.4.4.1 Varying environmental conditions

Studies reported on the production of halocarbons by microalgae by manipulating the environmental conditions in the laboratory. The parameters put to test include different light stress/irradiance (Moore *et al.*, 1996; Scarratt & Moore, 1999; Hughes *et al.*, 2006), photosynthesis activity (Hughes *et al.*, 2006) and elevated ozone levels (Thorenz *et al.*, 2014). Most of the studies on halocarbon emission by phytoplankton are mainly associated with the factor of different growth stages (Moore *et al.*, 1994; Tokarzcyk & Moore, 1994; Tait & Moore, 1995; Manley & de la Cuesta, 1997; Sæmundsdottir & Matrai, 1998; Scarratt & Moore, 1999).

(i) Growth phase

The amount of halocarbons emitted are largely measured in terms of which growth phase/stage the microalgae are in. The four growth stages in relation to halocarbon emission can be divided into lag phase, acceleration phase, stationary phase and death phase.

Tokarzcyk & Moore (1994) reported a suit of halocarbons, namely CHBr₃, CHBr₂Cl, CH₂Br₂ emitted at different amounts *by Porosira glacialis* and *Nitzschia sp.* (CCMP 580) at different growth phases under continuous cool white light with 12 μ mol m⁻² s⁻¹ photosynthetic active radiation at 6°C. The purge and trap system and gas chromatography equipped with electron capture detector were used to analyze and detect the amount of halocarbon emitted. A maximum of 2260 pmol L⁻¹ of CHBr₃ emitted by *Nitzschia* sp. (CCMP 580) was reported during its stationary phase (Day 17) at a cell density of 6 x10⁵ cell mL⁻¹ while 990pmol L⁻¹ of CHBr₃ emitted by *Porosira glacialis* was reported at a cell density of about 100 x10³ cell mL⁻¹ on the same day in stationary phase. An increase in emission rate of CHBr₃ in exponential growth phase from both

Nitzschia sp. (about 250- 2250 pmol L⁻¹) and *Porosira* culture (100- 800 pmol L⁻¹) was reported. The emission of CH₂Br₂ was also reported, reaching as high as 200 and 350 pmol L⁻¹ in *Nitzschia* sp. and *Porosira* cultures, respectively. Lower concentration of CH₂CII, about 16 pmol L⁻¹, emitted by *Nitzschia* sp. at a cell density of 600 x10³ cell mL⁻¹ was also reported.

Sæmundsdottir and Matrai (1998) reported a high emission rate of methyl bromide (CH₃Br), 30.1 pg CH₃Br µg⁻¹ Chl-a day⁻¹ by *Phaeocystis* sp. (CCMP 628) starting in late exponential or stationary phase, when cultured under 63 µEinstein s⁻¹ m⁻², 12:12 h light: dark cycle at 22°C. Other species such as *Amphidium carterae, Chaetoceros diversum, Hemiselmis rufescens, Pycnococcus provasolii, Pavloca* sp., *Prorocentrum micans* and *Chaetoceros atlanticus* produced 1.7- 9.3 pg CH₃Br µg⁻¹ Chl-a day⁻¹.

Scarratt and Moore (1999) reported the emission of CHCl₃ and CH₃I by the red microalga, *Porphyridium purpureum*. Of all the chlorinated hydrocarbon compounds, including C₂HCl₃, C₂Cl₄, CH₂Cl₂ and CHCl₃, over 4 to 10 days culture, only CHCl₃ was detected, with an emission rate range from 3.9 to 7.8 x10⁻⁷ mol g Chl-a⁻¹ day⁻¹. CH₃I was emitted by the same species and detected with an emission rate ranging from 4.8 x10⁻⁷ to 1.2×10^{-6} mol g Chl a⁻¹ day⁻¹.

A study by Hughes *et al.* (2011) reported that the changes in the emitted CH₃I concentrations are closely followed by increase in cell density in *Prochlorococcus marinus* (CCMP 2389). The amount of CH₃I emitted during the logarithmic phase (Day 1-5) is less than 100pmol L⁻¹. A maximum concentration of 690.5 pmol L⁻¹ of CH₃I was reached during mid- growth phase (Day 14). Manley & de la Cuesta (1997) reported a decrease in CH₃I production by *Nitzschia punctata* in the exponential phase where the number of cells in the culture peaked whilst an increase in the CH₃I (1.4 pmol of CH₃I) production occurred in the lag phase, but decreased to 8 pmol L⁻¹, after 16 days of growth.

In 2008, Colomb et al. screened several volatile organic compounds (VOCs) including CHCl₃, CH₃Cl, CH₃Br and CHBr₃, in *Emiliania huxleyi*, Calcidiscus leptoporus, Phaeodactylum tricornutum, Chaetoceros neogracilis and Dunaliella tertiolecta cultured under constant irradiance of 250 µEistein s⁻¹ m⁻², temperatures from 20 to 25°C adapting to a 12:12 h light:dark cycle. The sampling of the VOCs emissions were done in the middle of the light cycle during 4 days for Chaetoceros neogracilis and Emiliania huxleyi and 3 days for the rest of the taxa. Chaetoceros neogracilis and Phaeodactylum tricornutum were reported to emit CH₃Br the most, with 0.007 and 0.002 pmol L⁻¹/ Chla. Calcidiscus leptoporus was reported to emit 0.0038 pmol L^{-1} / Chl-a of CHBr₃. *Emiliania huxlevi* and *Chaetoceros neogracilis* emitted 0.002 pmol L⁻¹/ Chl-a of CHBr₃ whereas Dunaliella tertiolecta and Phaeodactylum tricornutum emitted 0.0001 pmol L⁻¹/ Chl-a of CHBr₃. Calcidiscus leptoporus was reported to emit approximately 1100 pptv of CHCl₃ while the rest reported to emit lower than 200 pptv. Colomb et al. (2008) also reported Calcidiscus leptoporus and Chaetoceros neogracilis to emit more than 2000 pptv of CH₃Cl, while *Phaeodactylum tricornutum* and *Dunaliella tertiolecta* reported to emit about 1750 pptv whereas *Emiliania huxylei* was reported to emit about 500 pptv of CH₃Cl.

Study has shown that the rate of CH_3I production in a monospecific *Synechococcus* culture can vary by at least an order of magnitude depending on physiological state (Hughes et al., 2011). One main factor that affects the emissions in different growth phases is the inoculation level. Morris et al. (2008) found that a high inoculation level (>3.5 x 10⁵ cell mL⁻¹) allows consistent growth in an axenic *P. marinus* culture and below this cells fail to growth well. Failure to grow well leads to cell stress and this negatively affects the cell growth in different growth phases and the emission rate of halocarbons, specifically CH_3I , into the surrounding medium (Morris et al., 2008). The emission of

other halogenated compounds such as brominated and chlorinated compounds, however, may not neccesarily be dependent on the cell growth.

(ii) Irradiance

The Sun has the ability to emit as high as approximately 2000 μ mol m⁻² s⁻¹ light radiating intensity to the Earth (Tibbitts, 1994). Nonetheless, the irradiance levels are subjected to fluctuation due to the constant changes in the weather. A few studies had put this factor to test to see if the changes of light intensity would influence the emission of halocarbons by the microalgae.

In 1996, Moore et al. reported a trend of higher concentrations of CH₂Br₂ and CHBr₂Cl produced in *Nitzschia* sp. (CCMP 580) and *Porosira glacialis* cultures with a higher level of illumination. In 40 μ mol quanta m⁻² s⁻¹ as low light, CH₂Br₂ emitted by *Nitzschia* sp. ranged from 0- 200 picoMolar (pM) while in higher light, the range of CH₃Br₃ emitted increased from 0 to 1300 pM over a period of 30 days. Nevertheless, this tendency was contrary to what was found by Hughes et al. (2006) and Scarratt & Moore (1996; 1999). Hughes et al. 2006 reported 2.5-2.8 pmol L⁻¹ of CH₃I emission over a period of 25 hours in both low light (47 μ mol m⁻² s⁻¹) and high light (250 μ mol m⁻² s⁻¹) for *Tetraselmis* sp. The same can be seen for *Emiliania huxylei* and *Thalassiosira pseudonana*, with CH₃I emission ranging 0.2-0.38 pmol L⁻¹ and 0.2- 0.3 pmol L⁻¹, respectively. Methyl halide production by three phytoplankton species (Phaeodactylum tricornutum, Phaeocystis sp., Thalassiosira weissflogii) was found to be not directly dependent on photosynthesis, but was closely related to biomass (Scarrat & Moore, 1996). Phaeocystis had the highest methyl halide production rate, while total CH₃Cl and CH₃Br increased higher during stationary phase of growth. Carbon limitation on the medium increased the production of halocarbons more than nitrogen limitation. Scarratt & Moore (1999) revealed that exposing the cultures of *Porphyridium purpureum* to high light intensity (800 µmol

quanta m⁻² s⁻¹) in excess of the acclimated irradiance level (20 μ mol quanta m⁻² s⁻¹) did not stimulate the production of the measured halocarbon compounds (CH₂Cl₂, C₂HCl₃, C₂Cl₄). In fact, the production of CH₃I dropped when exposed to high irradiance over a period of 24 hours. Scarratt & Moore's findings were supported by the work of Hughes *et al.* (2006) in a sense that light-induced stress does not induce the release of iodocarbon in any of the cultures examined.

(iii) Elevated ozone level

Thorenz *et al.* (2014) reported the only study of the influence of different elevated ozone level on the production of halocarbons by microalgae. To resemble the natural condition of the stratosphere, a glass chamber tube was continuously channelled with synthetic air flow at 3.4 L min⁻¹ over stirred algae suspension with ozone (100 ppb) and without ozone, each done separately. It was found that the release of several tested halocarbons, CHBr₃, CH₃I, CH₂CII and CH₂I₂, was not dependent on the high or low ozone level. This conclusion was made based on its insignificant differences in halocarbon emission rates, ranging 0.030- 0.098 ng min⁻¹ m⁻² for CH₃I, 0.003- 0.039 ng min⁻¹ m⁻² for CH₂CII, 0.073- 0.117 ng min⁻¹ m⁻² for CH₂I₂ and 0.503- 0.549 ng min⁻¹ m⁻² for CHBr₃ for *Porosira glacialis*. *Mediopyxis helysia* emitted a range of 9.90- 21.94 nmol L⁻¹ of iodide and 397- 499 nmol L⁻¹ of iodate, while *P. glacialis* emitted a range of 7.32- 19.71 ng min⁻¹ m⁻² of iodide and 408- 478 ng min⁻¹ m⁻² of iodate.

2.4.4.2 Halocarbon emissions and photosynthesis

The photosynthetic performance based on fluorometry was used by Hughes *et al.* (2006) to measure the maximum quantum yield (F_v/F_m) , which is an indicator of algal cell stress and the cell viability of the microalgae. Hughes *et al.* 2006 reported that under 14:10 h light:dark cycle at 15°C, there were no changes in the concentrations of CH₃I and CH₂CII relative to the control in any of the *Emiliania huxylei, Tetraselmis* sp. and

Thalassiosira pseudonana cultures, although the F_v/F_m in both control (low light at 47 µmol photons m⁻² s⁻¹) and high light (250 µmol photons m⁻² s⁻¹) treatments changed. In both the high and low light conditions, *Emiliania huxleyi* was reported to emit a range of 0.25 to 0.35 pmol L⁻¹ of CH₃I and approximately 2.5 to 4.5 pmol L⁻¹ of CH₂CII; *Tetraselmis* sp. was reported to emit about 2.5 to 2.8 pmol L⁻¹ of CH₃I and a range of 0.5 to 0.3 pmol L⁻¹ of CH₂CII; *Thalassiosira pseudonana* was reported to emit a range of 0.2 to 0.3 pmol L⁻¹ of CH₃I while no emission of CH₂CII was detected in both light conditions. The finding of a decrease in the fluorescence-based F_v/F_m for PSII (Photosystem II) due to high light stress revealed the oxidative damage to the photosynthetic apparatus in the cells. They concluded that iodocarbon release is not associated with light stress and is not likely to be in any connection with the protection against oxidative damage in the microalgae studied.

In another by Hughes *et al.* (2011), it was found that despite the decrease in F_v/F_m over a period of 22 days (from early exponential phase onwards), an increase in CH₃I concentration was observed throughout the 22 days experiment under an irradiance of 40 µEinstein m⁻² s⁻¹ with 14:10 h light:dark cycle at 22°C. During the increase of cell densities in exponential phase (from 4.0 x10⁷ cells mL⁻¹ (Day 2) to 8.0 x10⁷ cells mL⁻¹ (Day 7), *Prochlorococcus marinus* (CCMP 2389) showed a decrease in F_v/F_m of about 0.67 (Day 2) to 0.50 (Day 7) while an increase in CH₃I emission of approximately 20 pmol L⁻¹ to 100 pmol L⁻¹ in Day 2 and 7 respectively was reported. A decrease in F_v/F_m (0.5 to 0.3) on Day 7 (cell density at 8 x10⁷ cell mL⁻¹) to 16 (cell density of about 0.5 x10⁷ cell mL⁻¹) cultures showed an increase in the emission of CH₃I, from about 100 pmol L⁻¹ to 450 pmol L⁻¹. *Synechococcus* sp. (CCMP 2370) was reported to emit about 2.0 to 4.0 pmol L⁻¹ of CH₃I, which was close to the amount of CH₃I detected in the medium-only control.

CHAPTER 3: METHODOLOGY

3.1 Microalgal cultures

Three common tropical marine algal strains from the University of Malaya Algae Culture Collection (UMACC) were used for all experiments; the cyanophyte *Synechococcus* sp. UMACC 370 and the bacillariophyte *Amphora* sp. UMACC 370 were isolated from shrimp ponds connected to the Straits of Malacca in Kuala Selangor, Malaysia; while the chlorophyte *Parachlorella* sp. UMACC 245 was isolated from the east-coast waters facing the South China Sea in Terengganu, Malaysia. The strains represent three different classes of microalgae that are abundant in the local regions. Stock cultures were grown in Provasoli Medium (Prov50) (CCMP, 1996) under a 12h light:12h dark cycle and at a temperature of 25 ± 1 °C in an incubator shaker set at 100 rpm (PROTECH, model GC-1050). Silicate (Na₂SiO₃.9H₂O) was supplemented at 0.01g dm³ to the culture medium for *Amphora* sp. UMACC 370. The cultures were kept and maintained under axenic conditions using standard aseptic techniques; glassware and growth media were sterilized by autoclaving (15 min at 121°C) before use.

The isolated microalgae, *parachlorella* sp. UMACC 245, *synechococcus* sp. UMACC 371 and *amphora* sp. UMACC 370 as shown in respective Figures 3.1, 3.2 and 3.3 were imaged under Field Emission Scanning Microscope (FESEM) and light microscope for taxonomic identification. *Parachlorella* sp. UMACC 245 is unicellular, non-motile, coccoid, and has a size of about 3.5 - 4.0 µm in diameter. *Synechococcus* sp. UMACC 371 is unicellular, coccoid, non-motile and has a size of approximately 1.0 - 2.5 µm in diameter. *Amphora* sp. UMACC 370 is unicellular, non-flagellates and has a diameter size of about 10 - 11 µm. Based on the morphology and phylogenetic tree (Appendix A), the isolated *Amphora* sp. UMACC 370 appears to be most closely related to *Amphora Subtropica*, given 100% bootstrap support based on rcbL gene. However, the differences in nucleotide composition between the two sequences of *Amphora* and that the sequence

of *Amphora* sp. UMACC 370's actual identity may not be deposited in Genbank indicates that the isolated *Amphora* sp. UMACC 370 may be a new species.

A flow chart of research work is provied as shown in Figure 3.4.



Figure 3.1: *Parachlorella* sp. UMACC 245 under (a) FESEM using High vacuum mode (60 000x magnification) and (b) light microscope.



Figure 3.2: *Synechococcus* sp. UMACC 371 under (a) FESEM using High vacuum mode (30 000x magnification) and (b) light microscope.



Figure 3.3: *Amphora* sp. UMACC 370 under (a) FESEM using High vacuum mode (23 000x to 100 000x magnification) and (b) light microscope.

Figure 3.4: Flow-chart of research work



Thesis Writing

3.2 Experiment 1: Optimization of selected parameters for studies

3.2.1 Selection of suitable growth media

The objective was to select the suitable culture medium for the isolated microalgae. Three different growth media, Provasoli 50 (CCMP, 1996), F/2 (Guillard & Ryther, 1962) and Diatom Medium (Beakes *et al.*, 1988), were used to determine the most suitable medium for the growth of the three microalgae. The Optical Density (OD_{620nm}) of the cultures of the three microalgae were grown at 0.2 with 10% inoculum of exponential phase for a period of 12 days. The cultures of total volume of 150 mL in 250 mL conical flasks were incubated in the incubator shaker (100 rpm) at 25°C with irradiance of 40 µmol photons m⁻² s⁻¹ on a 12h light:12h dark cycle. OD_{620nm} readings were taken on a daily basis. This was done in triplicates.

3.2.2 Profiling algal growth

The objective was to profile the basic growth of the microalgae under laboratory conditions. The three local microalgal strains were grown in batch cultures of 150 ml in 250 mL conical flask with starting inoculum of 10% at OD_{620nm} 0.2 in Prov50 (CCMP, 1996) medium, with silicate (Na₂SiO₃.9H₂O) supplemented at 0.01g dm⁻³ to the culture medium for *Amphora* sp. UMACC 370, for a growth period of 14 days. The starting pH and salinity of the medium at the beginning of the experiment were set at 8 and 30 ppt, respectively. The cultures were grown under a 12h light:12h dark cycle and at 25 ± 1 °C in an incubator shaker (Model: GC-1050) under axenic conditions. Irradiance level in the growth chamber was maintained 40 ± 5 µmol photons m⁻² s⁻¹ for all the cultures.

The biomass of the cultures were determined every 2 days from the growth based on OD_{620nm} , chlorophyll-a (Vello et al., 2014; Strickland & Parsons, 1968) and cell count was done using Bright-field Neubauer haemacytometer (Marienfeld-Superior, Germany) under a light microscope (Vello *et al.*, 2014). Readings from other responding variables such as the carotenoids and pH were determined and specific growth rates (μ) were calculated. F_v/F_m , the maximum quantum yield, was also measured as an indication of cells's health state.

3.2.3 Determination of incubation time

The objective was to determine the air-tight incubation time based on cells' state of health (F_v/F_m) using PAM (Pulmonary Amplitude Modulation) Fluorometry (Hughes et al., 2011; Keng *et al.*, 2013). Cultures of the three microalgae were grown in triplicates for a total of eight hours, with F_v/F_m measured every hour. The value of the maximum quantum efficiency of photosystem II, denoted as F_v/F_m (where F_v is the variable fluorescence measured as the difference between maximum (F_m) and minimum (F_o) fluorescence in dark-adapted culture), was estimated using a Water PAM Fluorometer (Walz, Model: WATER-ED, S/N:EDEE0238 Germany). Samples from each culture were dark-adapted for 15 minutes prior to F_v/F_m determination.

3.3 Experiment 2: Determining suitable cell density for halocarbon studies

A short-term experiment to determine the suitable cell density in the cultures was conducted prior to the growth cycle experiment. The objective was to ensure that the cell density in the cultures during the growth cycle studies was sufficient to emit detectable levels of a suite of halocarbons by the GC-MS system used. The optical density at 620nm $(OD_{620 nm})$ of the cultures of the three microalgae were adjusted to 0.2, 0.3 and 0.4 for a growth period of four days, prior to measurement of the halocarbons. The cultures of total volume of 150 mL in 250 mL conical flasks were incubated in the incubator shaker (100

rpm) at 25°C with irradiance of 40 μ mol photons m⁻² s⁻¹ on a 12h light:12h dark cycle. The procedure for halocarbon determination is given below (Section 3.4.1).

3.4 Experiment 3: Effects of algal growth cycle on halocarbon emission

The objective of this experiment was to study how the emissions of halocarbons were affected by the difference in growth stages, specifically the exponential and stationary phase.

3.4.1 Experimental design

All three microalgal cultures were grown in batch culture with an inoculum size of 10% of a log phase culture, standardized at an optical density of 0.4 at 620nm (OD_{620nm}). The cultures of 150 mL total volume, were grown in 250 mL conical flasks in an incubator shaker (100 rpm) at 25°C with irradiance of 40 µmol photons m⁻² s⁻¹ on a 12h light:12h dark cycle. The experiments with each microalga was conducted in triplicates, with measurements done every two days. The experiment was completed after 12 days of growth. A control with culture medium but no microalgal inoculum was set up in triplicate for each microalga.

Every two days, 60 mL of cultures were removed from each triplicate flask and centrifuged (3000 rpm for 10 min) and replenished with fresh medium, then drawn out into a 100 mL glass syringe to be incubated air-tight for 4 hours. Incubation time was set to 4 hours to achieve sufficient amount of halocarbons trapped in the medium while minimizing cell stress. The amount of concentration for each halocarbon was obtained by subtracting the concentration of the sample to the control. Samples were also collected at the same time for biomass estimation using various parameters as described below (Section 3.4.5). This was to allow calculation of the emission rate through normalization of real-time biomass to the concentration of each compound emitted by microalgae over the culture period.

Emission rates were calculated by normalizing emission of halocarbons to the realtime biomass, both chlorophyll-a (pmol mg⁻¹ day⁻¹) and cell density (pmol cell⁻¹ day⁻¹). The formula to determine emission rate for this study is as follows:

$$Emission \ rate = \frac{Emission}{Biomass} \div (Incubation \ time \ x \ \frac{1 \ day}{24 \ hour})$$

where,

Emission = the concentration of halocarbon emitted at pmol L^{-1} Biomass = chlorophyll-a content (mg L^{-1}) or cell density (cell m L^{-1})

Incubation time = 4 hours

The state of the cells was determined using PAM Fluorometry (Hughes *et al.*, 2011; Keng *et al.*, 2013). The value of the maximum quantum efficiency of photosystem II, denoted as F_v/F_m (where F_v is the variable fluorescence measured as the difference between maximum (F_m) and minimum (F_o) fluorescence in dark-adapted culture), was estimated using a Water PAM (Pulmonary Amplitude Modulation) (Walz, Model: WATER-ED, S/N:EDEE0238 Germany) before and after the gas-tight incubation period to indicate the cells' health. Samples from each culture were dark-adapted for 15 minutes prior to F_v/F_m determination.

After 4 hours of incubation, the culture from the incubation syringes was gently swirled and extracted into 100 mL glass syringes through a two-syringe (0.2 μ m Merck filter unit) closed filter system to ensure no ingress of air into the syringe. Figure 3.4 shows the transfer of culture-filtered medium through an enclosed system.



Figure 3.5: Two-syringe collection system

The medium subsample left in the syringe was then ready for halocarbon analysis. The experiments were repeated for all three marine tropical microalgae.

All cultures were kept and maintained under axenic conditions using standard aseptic techniques; glassware and growth media were sterilized by autoclaving (15 min at 121°C) before use. Lysogenic broth (LB) agar plate was used to test and ensure the axenicity of the inoculum cultures.

3.4.2 Analysis for halocarbons

All halocarbon analyses were carried out using a purge-and-trap system developed by the University of East Anglia (UEA), UK (Hughes *et al.*, 2006) equipped with an Agilent Technologies 7890A gas chromatograph (GC). The GC was fitted with a J&W 60 m DB-VRX capillary column (film thickness 1.40 µm; internal diameter 0.25 mm).



Figure 3.6: Gas-Chromatography Mass-Spectrometry (left) and Purge-&-Trap System (right)

The extracted medium subsamples that had been injected into the system were purged for 15 minutes using oxygen-free nitrogen (OFN) at a flow rate of 40 mL min⁻¹. Any aerosols or particles in the bypassing purged gas would be removed through the stuffed glass wool held in a glass tubing. Water vapor in the bypassing of the purged gas was removed through a molecular sieve followed by a counter-flow Nafion dryer (Perma-Pure) using OFN at a rate of 100 mL min⁻¹. The targeted compounds were then trapped and cryogenically focused synchronously purging in a stainless-steel tubing coil immersed in liquid nitrogen at a temperature of -150°C, aided by a thermostated heating device for a total of 15 minutes.

Then to allow sample desorption, liquid nitrogen was quickly swapped with boiling water in a flow of high-purity Helium at 1 mL min⁻¹ via a heated (95°C) transfer line to the GC. As the run starts, the oven was initially held at 36°C for 5 min, followed by heating up to 200°C at 20°C min⁻¹, and lastly heated up to 240 °C at a rate of 40 °C min⁻¹. The quantification and identification of the compounds were determined by an Agilent 5975C mass-selective detector (MSD), operated in Single Ion Mode. Data was collected between 4 and 18 min. Calibrations for all compounds were done using gravimetrically

prepared liquid standards (Sigma-Aldrich) mixed in high-performance liquid chromatography-grade methanol (Fischer Scientific) injected into medium samples.

3.4.3 Calibration of halocarbon standards

Calibrations for all compounds (CH₃I, CHBr₃, CHCl₃, CH₂Br₂, CHBr₂Cl) were done using gravimetrically prepared liquid standards (Sigma-Aldrich) mixed in highperformance liquid chromatography-grade methanol (Fischer Scientific) injected into medium samples. The amount of halocarbon concentration from samples and phytoplankton-free controls were calculated based on a five-point calibration curve that plots concentration against peak area. The regression coefficient (r^2) for the linear calibration curve was above 0.95. Deuterated-iodomethane (CD₃I) (ARMAR chemicals) and deuterated-diiodomethane (CD_2I_2) (Sigma-Aldrich) were injected into every medium sample before the halocarbon analysis as a way to monitor and correct for drift in the detector sensitivity (Hughes et al., 2006). A loss of peak area from the internal standards due to the drift was corrected and equated to the original peak area as initially detected. Peak areas originated from analyte of interest, which was the halocarbons detected from the samples or controls, were also corrected following the same ratio as the surrogate standards did. The relative response, halocarbon concentration, was then obtained from the calibration that plots concentration against integrated peak area.

3.4.4 Detection limit and precision of the system

All halocarbon compounds were identified with their individual quantifying ions (Table 3.1) under the single ion mode (SIM) selected under the GCMS software programme run.

Compounds	Quantifying	Retention	Detection	Precision
	Ion (m/z)	time	Limit	(%)
		(min)	(pmol L ⁻¹)	(n=6)
CHBr ₃	173	15.61	0.30	10.3
CH ₃ I	142	8.18	0.20	5.9
CHCl ₃	83	10.33	0.50	7.3
CHBr ₂ Cl	129	14.09	0.05	9.8
CH_2Br_2	174	12.19	0.30	7.9

Table 3.1: Summary of halocarbons extracted by purge-and-trap and analyzed using GC-MSD and their associated quantifying ion, retention time, detection limit and precision.

3.4.5 Cell biomass determination

The biomass was determined based on cell number, which was counted using haemocytometer (Vello *et al.*, 2014). The chlorophyll-a content (chl-a) was determined by harvesting the microalgal cells by Millipore filtration using filter paper (Whatmann GF/C, 0.45 μ m). The chl-a of the microalgae were extracted using acetone and left overnight 4°C in the dark (Vello *et al.*, 2014; Strickland & Parsons, 1968). The absorption of the extract was measured at 665nm, 645nm and 630nm. Chl-a was calculated using the formula as follows:

$$Chl a (mg m^{-3}) = (Ca x Va)/Va$$

where, $Ca = 11.6 (OD_{665nm}) - 1.31(OD_{645nm}) - 0.14(OD_{630nm})$

Va = Volume of acetone (mL) used for extraction

Vc = Volume of culture (L)

$$Chl a (mg L^{-1}) = Chl a (mg m^{-3})/1000$$

The specific growth rate (ų, day⁻¹) for all cultures were based on calculated real-time biomass (chl-a and cell number) using the formula as follows:

$$y, day^{-1} = \frac{\ln(N_2/N_1)}{(t_2 - t_1)}$$

where N_2 , is OD_{620nm} at t_2 , N_1 , is real-time biomass at t_1 , and t_2 , t_1 are time periods within log phase (Strickland & Parsons, 1968).

To obtain dry weight, a pre-weighed glass fiber filter was used to filter 30.0 mL of culture. The filter paper was then dried in the oven at 100 °C for 24 hours, cooled in a dessicator before weighing. Dry weight (DW) was calculated as follows:

 $DW (mg L^{-1}) = \frac{(wt of filter paper \& algae) - wt of filter paper}{Volume of algal culture (L)}$

3.5 Experiment 4: Effects of varying irradiance on halocarbon emission

The objective of this experiment was to study how the emission of halocarbons and photosynthetic performance (F_v/F_m) of the microalgae were affected when exposed to a range of irradiance.

3.5.1 Experimental set-up

Three microalgal cultures, *Synechococcus* sp. UMACC 371, *Amphora* sp. UMACC 370 and *Parachlorella* sp. UMACC 245 were grown in batch culture with a starting inoculum size of 10% of a log phase culture, standardized at an optical density of 0.4 at OD_{620nm} . The cultures of 150 mL total volume, were grown in 250 mL conical flasks in an incubator shaker (100 rpm) at 25°C with irradiance of 40 µmol photons m⁻² s⁻¹ on a 12h light:12h dark cycle using F30T8/D HITACHI Fluorescent lamps 28W for up to 4 days to achieve exponential phase. On Day 4 during its light cycle, the cultures were exposed to 0, 40 (control) and 120 µmol photons m⁻² s⁻¹ for 12 hours. Light source was adjusted higher by adding the number of fluorescent tubes onto the incubator. Prior (t₀) and post (t₁) of the light-exposure period, samples were removed from each triplicate flasks and centrifuged (3000 rpm for 10 min) and replenished with fresh medium, drawn

out into a 100 mL glass syringe to be incubated air-tight for 4 hours. After 4 hours of incubation, the culture from the incubation syringes was gently swirled and extracted into 100 mL glass syringes through a two-syringe (0.2 μ m Merck filter unit) closed filter system to ensure no ingress of air into the syringe. The medium subsample left in the syringe was then ready for halocarbon analysis. Experiments were repeated for all three marine tropical microalgae.

3.5.2 Analysis and calibration of halocarbons

The procedure of analysis of halocarbon using GCMS equipped with Purge-and-Trap System was described in section 3.4.2.

3.5.3 Cell biomass determination

Samples were also collected at the same time for biomass estimation using various parameters, including chlorophyll-a and cell number as described in section 3.4.5. This was to allow calculation of the emission rate before and after 12hr of light-exposure through normalization of biomass to the emission of each compound emitted by microalgae.

3.5.4 Determination of photosynthetic parameter, F_v/F_m

The photosynthetic parameter, F_v/F_m was estimated by using a Water PAM (Pulmonary Amplitude Modulation) (Walz, Model: WATER-ED, S/N:EDEE0238 Germany) to indicate the changes of stress level before and after the light exposure, as well as before and after the gas-tight incubation period to indicate cells' health. Samples from each culture were dark-adapted for 15 minutes prior to F_v/F_m determination. A control with culture medium but no microalgal inoculum was set up in triplicate for each microalga.
3.6 Statistical Analysis

One-Way ANOVA was used to test the significance (p<0.05) of emission of all five compounds detected at early growth stage of the three microalgae at different OD_{620nm}.

Repeated Measures-ANOVA was used to test the significance (p<0.05) of emissions of all the five compounds by the three different microalgae within the culture period. Pearson Product-Moment correlation coefficient (r) was used to analyze the emission rate of the five detected compounds in term of chlorophyll-a, cell density and both.

Factorial-ANOVA was used to test the significance (p < 0.05) of the means of F_v/F_m through homogenous grouping for the three microalgae. Pearson Product-Moment correlation coefficient (r) was used to analyze the emission rates of the five compounds in terms of chlorophyll-a, cell density and both, as well as the correlation between F_v/F_m and halocarbon emission rates of the three taxa with irradiance. Pairwise comparison through Bonferroni adjustment was used to analyze the relationship between irradiance and halocarbon emission rates across the microalgae. Statistical analyses were done using the Statistica 8.0 and IBM SPSS Statistics software (p < 0.05). Data prior to data analyses were subject to the normality test using the skewness and kurtosis and results indicated that all data were distributed normally.

CHAPTER 4: RESULTS

4.1 Experiment 1: Optimization for halocarbon studies

4.1.1 Growth curves of microalgae in different culture media

Figure 4.1 show the growth curves of three tropical marine microalgae, (a) *Parachlorella* sp. UMACC 245, (b) *Synechococcus* sp. UMACC 371 and (c) *Amphora* sp. UMACC 370 grown in three media, namely f/2, Diatom (DM) and Prosavoli 50 (Prov 50) medium at OD_{620nm} 0.2. The three microalgae showed clear trends of higher biomass when grown in Prov50 based on OD_{620nm} as compared to the other two media, though the trend of growth under f/2 medium came close with Prov50 for *Parachlorella* sp. UMACC 245 and *Synechococcus* sp. UMACC 371. There were more clumpings observed in DM in comparisons with Prov50 and f/2 as the diatom cultures proceeded from exponential to stationary phase.





Figure 4.1: Growth curves based on Optical Density (OD_{620nm}) of three tropical marine microalgae, (a) *Parachlorella* sp. UMACC 245, (b) *Synechococcus* sp. UMACC 371 and (c) *Amphora* sp. UMACC 370 under different growth media over a period of 12 days. n = 3

Specific growth rates, ų (day⁻¹) of the microalgae under all three different growth media were calculated and summarized in Table 4.1. Based on the results, all microalgae had the highest specific growth rate when grown under Prov50 medium as compared to other media, F/2 and Diatom media. Hence, Prov50 medium was selected as the most suitable medium for all microalga culture.

	Parachorella	Synechococcus	Amphora sp.
Medium	sp. UMACC 245	sp. UMACC 371	UMACC 370
F/2	0.657	1.019	0.288
	(±0.034)	(±0.001)	(±0.008)
Diatom	0.685	1.159	0.189
	(±0.020)	(±0.042)	(±0.026)
Provasoli	0.705	1.163	0.320
50	(±0.019)	(± 0.022)	(±0.012)

Table 4.1: Specific growth rate, μ (day⁻¹) of three tropical marine microalgae based on exponential phase under different growth media. n = 3.

4.1.2 Basic growth profile of the selected microalgae

Three selected tropical marine microalgae were grown for 14 days to obtain respective growth curve profiles, from lag phase to stationary phase, through the measurements of chlorophyll-a (Figure 4.2), cell density (Figure 4.3), optical density (Figure 4.4) and carotenoids (Figure 4.5). Maximum quantum yield, F_v/F_m , for all three microalgae were recorded as shown in Figure 4.6 to be observed as cell stress indicator.



Figure 4.2: Growth curves of three tropical marine microalgae over a growth period of 14 days determined by chlorophyll-a. n = 3



Figure 4.3: Growth curves of three tropical marine microalgae over a growth period of 14 days determined by cell density. n = 3



Figure 4.4: Growth curve of three tropical marine microalgae over a growth period of 14 days determined by Optical Density (OD_{620nm}). n = 3



Figure 4.5: Carotenoids of three tropical microalgae over a growth period of 14 days. n = 3



Figure 4.6: Maximum quantum yield, F_v/F_m of three tropical marine microalgae over a growth period of 14 days. n = 3



Figure 4.7: pH of three tropical marine microalgae over a growth period of 14 days. n=3

Parachlorella sp. UMACC 245 and *Synechococcus* sp. UMACC 371 showed consistent trends of higher biomass in exponential phase as compared to *Amphora* sp. UMACC 370 based on chlorophyll-a, cell density, OD_{620nm} and carotenoids. Over the 14-day of culture period, tha range of F_v/F_m falls between 0.5-0.7, 0.3-0.4 and 0.5-0.6 for *Parachlorella* sp. UMACC 245, *Synechococcus* sp. UMACC 371 and *Amphora* sp. UMACC 370 respectively. The starting pH for the culture on Day 0 was 8.0. Over the 14 day growth period, the pH trend for *Parachlorella* sp. UMACC 245 and *Synechococcus* sp. UMACC 245 and *Synechococcus* sp. UMACC 245 and *Synechococcus* sp. UMACC 371 increased but decreased for *Amphora* sp. UMACC 370 during the exponential phase.

4.1.3 Selection of suitable air-tight incubation hours

The aim of this test was to determine the most suitable air-tight incubation time for the three microalgae so as to maximize the production of halocarbons while minimizing cells' physiological stress. Based on the F_v/F_m results as shown in Table 4.2, there was an indication of stress from the decrease of F_v/F_m starting 5th hour onwards for all three microalgae. Hence, 4 hours was the longest possible period for air-tight incubation. Whilst the F_v/F_m were slightly higher on the first three hours as compared to the 4th hour across all three taxa, 4 hours was ultimately selected as the most suitable incubation time

as one of the two aims in this test was to trap more halocarbons within the fixed incubation time. The cells were generally under non-stress condition after 4 hours of air-tight process.

		Maximum Quantum Efficiency (F _v /F _m), n=3			
	Air-Tight	Parachlorella sp.	Synechococcus sp.	Amphora sp.	
Time	Hours	UMACC 245	UMACC 371	UMACC 370	
9.50am	0 (control)	0.62 (±0.00)	0.38 (±0.01)	0.63 (±0.00)	
10.50am	1	0.63 (±0.00)	0.36 (±0.00)	0.62 (±0.00)	
11.50am	2	0.61 (±0.00)	0.37 (±0.01)	0.62 (±0.00)	
12.50am	3	0.61 (±0.00)	0.38 (±0.00)	0.60 (±0.00)	
1.50pm	4	0.61 (±0.00)	0.37 (±0.00)	0.60 (±0.00)	
2.50pm	5	0.59 (±0.00)	0.32 (±0.01)	0.59 (±0.00)	
3.50pm	6	0.60 (±0.00)	0.29 (±0.01)	0.56 (±0.00)	
4.50pm	7	0.60 (±0.00)	0.29 (±0.01)	0.52 (±0.00)	
5.50pm	8	0.59 (±0.01)	0.27 (±0.01)	0.49 (±0.01)	

Table 4.2: Comparison of F_v/F_m across 8 hours of air-tight incubation for *Parachlorella* sp. UMACC 245, *Synechococcus* sp. UMACC 371 and *Amphora* sp. UMACC 370.

4.2 Experiment 2: Emission of halocarbons at different cell densities

Only five halocarbons were detected in the emissions from the three microalgae. Table 4.3 shows the concentration (pmol L⁻¹) of halocarbons emitted by the three microalgae of different cell densities (OD_{620nm})*. a) *Synechococcus* UMACC 371; b) *Amphora* UMACC 370; c) *Parachlorella* UMACC 245. n= 3.

Synechococcus sp. UMACC 371 showed no emission at OD_{620nm} 0.2 for all five compounds except CH₃I and CHBr₃ on Day 4. At OD_{620nm} 0.3, all compounds showed no emission except CH₃I on Day 0, 2 and 4, and CHBr₃ on Day 4. At OD_{620nm} 0.4, emissions for all compounds were detected except on Day 0 and 4 for CHCl₃, Day 0 and 4 for CH₂Br₂ and Day 0 and 2 for CHBr₂Cl. OD_{620nm} 0.4 produced the highest (*p*<0.05) emission for all five compounds on Day 0, 2 and 4 as compared to OD_{620nm} 0.2 and 0.3. The inoculum of OD_{620nm} 0.4 was selected to be the starting inoculum for halocarbon emission study for *Synechococcus sp.* UMACC 371.

At OD_{620nm} of 0.2, there was no emission from *Amphora* sp. UMACC 370 except for CH_2Br_2 on Day 4 and CH_3I . At OD_{620nm} 0.3, emissions of all compounds were detected but were inconsistent from Day 0 to 4 as no emission was observed in some compounds. Higher emission (at least 1x and above) was detected at OD_{620nm} 0.4 across all five compounds from Day 0 to 4, except for $CHBr_2CI$ at Day 0. OD_{620nm} 0.4 produced the highest (p<0.05) emission compared to OD_{620nm} 0.2 and 0.3. The inoculum of OD_{620nm} 0.4 was selected for *Amphora* sp. UMACC 370 as the starting inoculum.

Parachlorella sp. UMACC 245 showed no emission at OD_{620nm} 0.2 for all five compounds across the 4-day experiment period and no emission at OD_{620nm} 0.3 except for two brominated compounds, CHBr₃ on Day 2 and CH₂Br₂ on Day 4. Emissions were detected for all compounds at OD_{620nm} 0.4, but not detected from Day 0 to 4. The inoculum of OD_{620nm} 0.4 produced the highest (*p*<0.05) emission amongst the three OD_{620nm} tested, and was selected to be the starting inoculum for the halocarbon emission studies. Tables 4.3: Emission of five halocarbons emitted by cultures of different OD_{620nm}. Concentration (pmol L⁻¹) of halocarbons emitted by three microalgae of different cell densities (OD_{620nm})*. a) *Synechococcus* UMACC 371; b) *Amphora* UMACC 370; c) *Parachlorella* UMACC 245. n= 3. Different letters denote standard deviation (SD) homogenous group (p<0.05) according to post-hoc Tukey's test.

		(a)	
CH3I	0.2*	0.3*	0.4*
Day 0	0.00^{b}	0.18 (±0.03) ^b	1.02 (±0.23) ^a
Day 2	0.00^{b}	$0.04 \ (\pm 0.03)^{\rm b}$	$0.11 \ (\pm 0.02)^{\rm b}$
Day 4	0.20 (±0.10) ^b	$0.07 \ (\pm 0.03)^{\rm b}$	0.18 (±0.06) ^b
CHCl3	0.2*	0.3*	0.4*
Day 0	0.00^{b}	0.00^{b}	0.00^{b}
Day 2	0.00^{b}	0.00^{b}	0.00 ^b
Day 4	0.00^{b}	0.00 ^b	1.03 (±0.19) ^a
CHBr ₃	0.2*	0.3*	0.4*
Day 0	0.00 ^c	0.00 ^c	0.02 (±0.01) ^b
Day 2	0.00^{c}	$0.00^{\rm c}$	$0.01 \ (\pm 0.00)^{b,c}$
Day 4	$0.03(\pm 0.00)^{b}$	$0.02 (\pm 0.00)^{b}$	0.05 (±0.01) ^a
CH2Br2	0.2*	0.3*	0.4*
Day 0	0.00^{b}	0.00^{b}	0.00^{b}
Day 2	0.00^{b}	0.00 ^b	$0.04 \ (\pm 0.03)^{a}$
Day 4	0.00^{b}	0.00 ^b	0.00^{b}
CHBr2Cl	0.2*	0.3*	0.4*
Day 0	0.00 ^b	0.00 ^b	0.00 ^b
Day 2	0.00^{b}	0.00^{b}	0.00^{b}
Day 4	0.00 ^b	0.00^{b}	0.01 (±0.00) ^a

		(b)	
CH ₃ I	0.2*	0.3*	0.4*
Day 0	$0.09 \ (\pm 0.04)^{d}$	$0.28 \ (\pm 0.03)^{d}$	$0.67 \ (\pm 0.10)^{b,c}$
Day 2	$0.12 (\pm 0.04)^{d}$	$0.27 \ (\pm 0.06)^{d}$	$0.87 \ (\pm 0.06)^{a,b}$
Day 4	$0.24 \ (\pm 0.04)^{d}$	$0.56 \ (\pm 0.12)^{c}$	$1.07 \ (\pm 0.08)^{a}$
CHCl3	0.2*	0.3*	0.4*
Day 0	0.00 ^d	$0.28 \ (\pm 0.13)^{b,c}$	$0.78 \ (\pm 0.14)^{a}$
Day 2	0.00^{d}	0.00 ^d	$0.37 \ (\pm 0.08)^{\rm b}$
Day 4	0.00^{d}	0.00^{d}	$0.11 \ (\pm 0.03)^{c,d}$
CHBr3	0.2*	0.3*	0.4*
Day 0	0.00 ^c	0.00 ^c	$0.15 (\pm 0.04)^{b}$
Day 2	0.00 ^c	0.00 ^c	$0.09 \ (\pm 0.01)^{\rm b}$
Day 4	0.00 ^c	$0.02 \ (\pm 0.00)^{c}$	$0.22 \ (\pm 0.06)^{a}$
CH2Br2	0.2*	0.3*	0.4*
Day 0	0.00 ^b	0.00 ^b	0.03 (±0.02) ^b
Day 2	0.00 ^b	$0.01 \ (\pm 0.00)^{\rm b}$	$0.14 \ (\pm 0.06)^{a}$
Day 4	0.00^{b}	$0.02 \ (\pm 0.01)^{b}$	$0.04 \ (\pm 0.02)^{\rm b}$
CHBr2Cl	0.2*	0.3*	0.4*
Day 0	0.00 ^c	0.00^{c}	0.00 ^c
Day 2	0.00 ^c	0.00 ^c	$0.02 \ (\pm 0.00)^{a}$
Day 4	0.00 ^c	$0.01 \ (\pm 0.00)^{a,b}$	$0.01 \ (\pm 0.00)^{\rm b}$

		(c)	
CH3I	0.2*	0.3*	0.4*
Day 0	0.00 ^c	0.00 ^c	0.12 (±0.03) ^a
Day 2	0.00 ^c	0.00°	$0.05 \ (\pm 0.03)^{\rm b}$
Day 4	0.00 ^c	0.00 ^c	0.00 ^c
CHCl3	0.2*	0.3*	0.4*
Day 0	0.00 ^b	0.00^{b}	$0.36 \ (\pm 0.17)^{a}$
Day 2	0.00 ^b	0.00^{b}	0.00^{b}
Day 4	0.00 ^b	0.00^{b}	0.00^{b}
CHBr ₃	0.2*	0.3*	0.4*
Day 0	0.00 ^c	0.00 ^c	0.00 ^c
Day 2	0.00 ^c	0.01 (±0.03) ^b	0.00 ^c
Day 4	0.00 ^c	0.00 ^c	$0.01 \ (\pm 0.00)^{a}$
CH2Br2	0.2*	0.3*	0.4*
Day 0	0.00 ^c	0.00 ^c	0.00 ^c
Day 2	0.00 ^c	0.00 ^c	$0.03 \ (\pm 0.01)^{a}$
Day 4	0.00 ^c	0.01 (±0.00) ^b	$0.01 \ (\pm 0.00)^{b}$
CHBr2Cl	0.2*	0.3*	0.4*
Day 0	0.00 ^b	0.00 ^b	0.00 ^b
Day 2	0.00 ^b	0.00 ^b	0.00 ^b
Day 4	0.00^{b}	0.00^{b}	$0.02 \ (\pm 0.01)^{a}$

4.3 Experiment 3: Halocarbon emission at different life-cycle stage

4.3.1 Growth curves of microalgae

The growth curves in terms of chlorophyll-a are shown in Figure 4.8 (a-c) and cell density as shown in Figure 4.9 (a-c), indicating the exponential and stationary phases for all three taxa (Table 4.4), and allowing the calculation of the specific growth rates (Table 4.5).



Figure 4.8: Growth curves based on chlorophyll-a. Cell growth phases of three tropical marine microalgae, (a) *Synechococcus* sp. UMACC 371; (b) *Parachlorella* sp. UMACC 245; (c) *Amphora* sp. UMACC 370 based on real-time biomass, chlorophyll-a (mg L⁻¹) over 12 days of culture period. n = 3



Figure 4.9: Growth curves based on cell density. Cell growth phases of three tropical marine microalgae, (a) *Synechococcus* sp. UMACC 371; (b) *Parachlorella* sp. UMACC 245; (c) *Amphora* sp. UMACC 370 based on real-time biomass, cell density (cell mL⁻¹) over 12 days of culture period. n = 3

Table 4.4: Algal growth stages determined by chlorophyll-a and cell density.
Selected range and representative points of exponential and stationary phases for the
three tropical marine microalgae are shown.

Taxa	Exponential phase		Stationary phase	
	Phase	Representative	Phase	Representative
	range	point	range	point
Synechococcus	Day		Day	
sp. UMACC 371	0—4		4—12	
Parachlorella	Day		Day	
sp. UMACC 245	0—4		4—12	
Amphora sp.	Day	Day 4	Day	Day 8
UMACC 370	0—6		6—12	
	Day			
	2 - 6 #			

For Amphora, the exponential phase ranged from day 2 to day 6.

Table 4.5: Specific growth rate. The mean of specific growth rate, ψ (day⁻¹) of the three tropical marine microalgae based on their exponential growth phase of chlorophyll-a and cell density. n = 3.

	Specific Growth Rate (ų)			
Taxa	Chlorophyll-a	Cell density		
Synechococcus sp. UMACC 371	0.66 (±0.01)	0.36 (±0.04)		
Parachlorella sp. UMACC 245	0.54 (±0.06)	0.64 (±0.07)		
Amphora sp. UMACC 370	0.27 (±0.04)	0.74 (±0.05)		

4.3.2 Photosynthetic performance as an indication of cells' health state

Figure 4.10 (a-c) show the maximum quantum yield (F_v/F_m) of three tropical marine microalgae across a period of 12 days before and after the 4-hour air-tight incubation. F_v/F_m values shown prior to air-tight incubation act as control level. The smallest difference in F_v/F_m before and after air-tight incubation ensured the production of halocarbons trapped during the incubation from cell culture was maximized while the cells remained healthy or minimally affected by the physiological stress created from an air-tight environment. Under ambient laboratory conditions, the healthy range of F_v/F_m for *Synechococcus* sp. UMACC 371, *Parachlorella* sp.245 and *Amphora* sp.370 were

within 0.3-0.4, 0.5-0.7 and 0.5-0.7 respectively. In general, the cells for all cultures were in the healthy F_v/F_m range. Hence, the emission of halocarbons were not under the influence of cell stress from the air-tight incubation.



Figure 4.10: Maximal quantum efficiency, F_v/F_m . The mean of F_v/F_m for (a) *Synechococcus* sp. UMACC 371; (b) *Parachlorella* sp. UMACC 245; (c) *Amphora* sp. UMACC 370 before and after incubation over 12-day culture period. n = 3.

4.3.3 Emission of halocarbons

Figure 4.11 (a-e) show the changes of concentration of the five halocarbons from the three microalgae cultures and Prov50 medium over the 12 experimental days. The halocarbon concentrations above the control level were the emission while concentrations below the control level were the consumption or loss of the halocarbons. In case of present study, the aim focused on the emission of the halocarbons from the three microalgae.











Figure 4.11: Changes of concentration of halocarbons detected from the three microalgae and Prov50 medium (controls) over a growth period of 12 days for compound (a) CHBr₃, (b) CH₃I, (c) CHCl₃, (d) CHBr₂Cl, and (e) CH₂Br₂. n = 3.

To obtain the amount of halocarbons emitted, the concentration of halocarbons from the Prov50 medium (controls) were subtracted from the concentration of halocarbons emitted from the microalgae. Figure 4.12 (a-e) show the concentrations of the five emitted halocarbons from the three tropical marine microalgae. The concentration of CH₃I was the highest at 1.6 pmol L⁻¹ from both stationary and exponential phases. *Amphora* sp. UMACC 370 showed significantly (p<0.05) higher concentration of CHBr₃, CH₃I, CHBr₂Cl during the stationary phase and CH₂Br₂ during the late exponential phase. *Synechococcus* sp. UMACC 371 in general emited higher amount of CH₃I during both exponential and stationary phases, as compared to other detected compounds. *Parachlorella* sp. UMACC 245 emitted low concentration across all compounds except CHBr₃ and CH₂Br₂ in both exponential and stationary phases and CH₃I during the exponential phase.











Figure 4.12: Emission of short-lived halocarbons. Concentration of halocarbon emitted by the three tropical marine microalgae across 12 experimental days for compound (a) CHBr₃, (b) CH₃I, (c) CHCl₃, (d) CHBr₂Cl and (e) CH₂Br₂. Bar charts which contain different alphabets denote significant difference at (p < 0.05). n = 3

4.3.4 Emission rates of halocarbons

The concentration of the five detected halocarbons were used to normalize to chlorophyll-a (Figure 4.13) and cell density (Figure 4.14) to determine emission rate. Both chlorophyll-a and cell density normalized emission rates for all five compounds across the three microalgae were in good agreement. *Amphora* sp. UMACC 370 showed significantly (p<0.05) higher emission rate of CH₃I, CHCl₃ and CH₂Br₂ in the exponential phase. In addition, emission rates of CHBr3 and CHBr2Cl were significantly higher (p<0.05) than other halogenated compounds during exponential and stationary phase. The emission rates of all five compounds for *Synechococcus* sp. UMACC 371 and *Parachlorella* sp. UMACC 245 were significantly (p<0.05) lower as compared to *Amphora* sp. UMACC 370.













Figure 4.13: Emission rate normalized to chlorophyll-a. Concentration of compound (a) CHBr₃, (b) CH₃I, (c) CHCl₃, (d) CHBr₂Cl and (e) CH₂Br₂ normalized to real-time chlorophyll-a for the three tropical microalgae across 12 experimental days. Bar charts which contain different alphabets denote significant difference at (p < 0.05). n =3

4.3.4.2 Normalization to cell density









Figure 4.14: Emission rate normalized to cell density. Concentration of compound (a) CHBr₃, (b) CH₃I, (c) CHCl₃, (d) CHBr₂Cl and (e) CH₂Br₂ normalized to cell density for the three tropical microalgae across 12 experimental days. Bar charts which contain different alphabets denote significant difference at (p < 0.05). n = 3

4.3.5 Comparison of emission rates of halocarbons by growth phases

Table 4.6 (a-c) show the estimated (upper and lower limits) emission rate of measured halocarbons under conditions of the experiments by the three tropical marine microalgae. *Amphora* sp. UMACC 370 showed the highest emission rates for methyl iodide (CH₃I) in both exponential and stationary phases, reporting 14.18 - 86.79 pmol (mg chla)⁻¹ day⁻¹ and 10.02 - 18.08 pmol (mg chla)⁻¹ day⁻¹ respectively when normalized to chlorophyll-a, and 2.05 - 24.05 pmol (10^9 cell)⁻¹ day⁻¹ (exponential) and 1.29 - 3.16 pmol (10^9 cell)⁻¹ day⁻¹ (stationary) when normalized to cell density, as compared to *Synechococccus* sp. UMACC 370 and *Parachlorella sp.* UMACC 245. Estimated emission rates of CH₃I for *Amphora* sp. UMACC 370, *Synechococccus* sp. UMACC 371 and *Parachlorella* sp. UMACC 245 based on chlorophyll-a and cell density were higher in exponential phase than in stationary phase, except in case of *Synechococcus* sp. UMACC 371 where the emission of CH₃I was lower in exponential phase as compared to its stationary phase.

Table 4.6: Emission rate at different growth phases. Concentrations of five halocarbons normalized to chlorophyll-a (pmol mg chl-a⁻¹ day⁻¹) and cell density (pmol (10⁹ cell)⁻¹ day⁻¹) at exponential and stationary phase for (a) *Synechococcus* sp. UMACC 371, (b) *Parachlorella* sp. UMACC 245 and (c) *Amphora* sp. UMACC 370.

(a)						
	Exponential phase		Stationary Phase			
Compound	pmol (mg	pmol (10 ⁹	pmol (mg	pmol (10 ⁹		
	chla) ⁻¹ day ⁻¹	cell) ⁻¹ day ⁻¹	chla) ⁻¹ day ⁻¹	cell) ⁻¹ day ⁻¹		
CHBr ₃	0.00 - 5.97	0.00 - 1.18	0.00 - 1.58	0.00 - 0.32		
CH ₃ I	0.00 - 12.27	0.00 - 2.70	0.74 - 2.23	0.16 - 0.79		
CHCl ₃	0.00 - 30.96	0.00 - 5.95	0.00 - 0.37	0.00 - 0.07		
CHBr ₂ Cl	0.00 0.13	0.00 - 0.07	0.00 - 0.07	0.00 - 0.01		
CH_2Br_2	0.00 - 8.23	0.00 - 1.58	0.00 - 0.21	0.00 - 0.04		

(b)					
	Exponential phase		Stationary Phase		
Compound	pmol (mg	pmol (10 ⁹	pmol (mg	pmol (10 ⁹	
_	chla) ⁻¹ day ⁻¹	cell) ⁻¹ day ⁻¹	chla) ⁻¹ day ⁻¹	cell) ⁻¹ day ⁻¹	
CHBr ₃	0.00 - 1.16	0.00 - 0.30	0.00 - 1.28	0.00 - 0.19	
CH ₃ I	0.00 - 3.36	0.00 - 0.83	0.00 - 1.02	0.00 - 0.23	
CHCl ₃	0.00 - 48.68	0.00 - 12.11	0.00 - 0.26	0.00 - 0.05	
CHBr ₂ Cl	0.00 - 0.22	0.00 - 0.05	0.00 - 0.08	0.00 - 0.01	
CH_2Br_2	0.00 - 2.63	0.00 - 0.66	0.00 - 0.33	0.00 - 0.04	

(c)					
	Exponential phase		Stationary Phase		
Compound	pmol (mg	pmol (10 ⁹	pmol (mg	pmol (10 ⁹	
	chla) ⁻¹ day ⁻¹	cell) ⁻¹ day ⁻¹	chla) ⁻¹ day ⁻¹	cell) ⁻¹ day ⁻¹	
CHBr ₃	0.00 - 22.46	0.00 - 5.97	0.45 - 8.81	0.09 – 1.59	
CH ₃ I	14.18 - 86.79	2.05 - 24.05	10.02 - 18.08	1.29 - 3.16	
CHCl ₃	0.00 - 48.51	0.00 - 12.90	0.00 - 1.27	0.00 - 0.15	
CHBr ₂ Cl	0.00 - 1.84	0.00 - 0.49	0.00 - 1.89	0.00 - 0.21	
CH ₂ Br ₂	0.00 - 14.04	0.00 - 5.85	0.00 - 2.77	0.00 - 0.44	

The estimated emission rates of CHBr₃, CH₂Br₂, CHCl₃, CH₃I and CHBr₂Cl for *Amphora* sp. UMACC 370 were all higher in exponential phase than in stationary phase, except the emission rate based on chlorophyll-a for CHBr₂Cl during exponential phase (1.84 pmol (mg chla)⁻¹ day⁻¹) was lower than in stationary phase (1.89 pmol (mg chla)⁻¹ day⁻¹). *Synechococcus* sp. UMACC 370 reported higher range of CH₃I and CHCl₃ emission rates in log phase than in stationary phase based on chlorophyll-a, whereas higher range of emission rates for CH₃I, CHBr₃, CHCl₃ and CH₂Br₂ based on cell density. *Parachlorella* sp. UMACC 245 reported lower estimated emission rates for CHBr₃, CHCl₃, CH₂Br₂ and CHBr₂Cl during exponential phase as compared to its stationary phase.

Estimated emission rate of CHCl₃ was higher in exponential phase as compared to stationary phase; 30.96 pmol (mg chla)⁻¹ day⁻¹ and 0.37 pmol (mg chla)⁻¹ day⁻¹ respectively for *Synechococcus* sp. UMACC 371, 48.51 pmol (mg chla)⁻¹ day⁻¹ and 1.27 pmol (mg chla)⁻¹ day⁻¹, respectively for *Amphora* sp. UMACC 370. *Parachlorella* sp. UMACC 245 had higher emission rates during the exponential phase as compared to its stationary phase based on chlorophyll-a and cell density.

Out of the three brominated compounds, estimated emission rates for CHBr₃ was higher than CH₂Br₂ and CHBr₂Cl during stationary phase based on chlorophyll-a across all three tropical marine microalgae. The estimated emission rates of CHBr₃, CH₂Br₂ and CHBr₂Cl by *Synechococcus* sp. UMACC 371 based on chlorophyll-a and cell density during exponential phase were higher as compared to their stationary phase. Higher estimated emission rates based on chlorophyll-a during stationary phase than in exponential phase was observed for CHBr₃ and CHBr₂Cl both by *Parachlorella* sp. UMACC 245 and *Amphora* sp. UMACC 370, except for CH₂Br₂ where emission rate during exponential phase was higher than its stationary phase. *Amphora* sp. UMACC 370 had at least two times higher of CH₂Br₂, CHBr₃ and CHBr₂Cl emission rates during both exponential and stationary phases based on chlorophyll-a. *Chlorella* sp. UMACC 245 showed the least emission rates of all three brominated compounds during exponential phase based on chlorophyll-a. *Chlorella* sp. UMACC 245 showed the least emission rates of all three brominated compounds during exponential phase based on chlorophyll-a. *Chlorella* sp. UMACC 245 showed the least emission rates of all three brominated compounds during exponential phase based on chlorophyll-a. *Chlorella* sp. UMACC 245 showed the least emission rates of all three brominated compounds during exponential phase based on chlorophyll-a. *Chlorella* sp. UMACC 245 showed the least emission rates of all three brominated compounds during exponential phase based on chlorophyll-a for the three brominated compounds during stationary phase than exponential phase was observed.





Figure 4.15: Total emission rate in percentage. Total rate of emission (%) of every five halocarbons in comparison amongst the three tropical marine microalgae based on (a) cell number and (b) chlorophyll-a.

Data for emission rate for all 12 days of cultures were combined (n=21) and were expressed and compared in percentage in terms of five compounds amongst the three microalgae. In Figure 4.15, *Amphora* sp. UMACC 370, consisting the majority of emission rate percentage for all five compounds, showed higher emission rate percentage as compared to *Synechococcus* sp. UMACC 371 and *Parachlorella* sp. UMACC 245. In

other words, *Amphora* sp. UMACC 370 was clearly a stronger emitter of the five halogenated compounds as compared to the other two taxa based on chlorophyll-a and cell density.

4.3.7 Correlation of detected halocarbons

The emission rates for all five compounds based on chlorophyll-a, cell density and both as summarized in Table 4.7 (a-c) were highly (p<0.01) correlated.

Table 4.7: Correlation of the halocarbons. Pearson Product-Moment correlation coefficient (r) of the emission rate from the five detected compounds in term of (a) chlorophyll-a, (b) cell density, (c) chlorophyll-a and cell density.

CHBr ₃	CH ₃ I	CHCl ₃	CHBr ₂ Cl	CH ₂ Br ₂
1.0000	0.7122**	0.4224**	0.6016**	0.4642**
0.7122**	1.0000	0.4828**	0.6390**	0.6195**
0.4224**	0.4828**	1.0000	0.3081*	0.6543**
0.6016**	0.6390**	0.3081*	1.0000	0.4659**
0.4642**	0.6195*	0.6543**	0.4659**	1.0000
	CHBr ₃ 1.0000 0.7122** 0.4224** 0.6016** 0.4642**	CHBr3CH3I1.00000.7122**0.7122**1.00000.4224**0.4828**0.6016**0.6390**0.4642**0.6195*	CHBr3CH3ICHCl31.00000.7122**0.4224**0.7122**1.00000.4828**0.4224**0.4828**1.00000.6016**0.6390**0.3081*0.4642**0.6195*0.6543**	CHBr3CH3ICHCl3CHBr2Cl1.00000.7122**0.4224**0.6016**0.7122**1.00000.4828**0.6390**0.4224**0.4828**1.00000.3081*0.6016**0.6390**0.3081*1.00000.4642**0.6195*0.6543**0.4659**

Number of replicates (n) = 63, ** = (p) < 0.01; * = (p) < 0.05

(b)	CHBr ₃	CH ₃ I	CHCl ₃	CHBr ₂ Cl	CH ₂ Br ₂
CHBr ₃	1.0000	0.7864**	0.6176**	0.8391**	0.6266**
CH ₃ I	0.7864**	1.0000	0.5964**	0.8489**	0.6430**
CHCl ₃	0.6176**	0.5964**	1.000	0.5872**	0.6872**
CHBr ₂ Cl	0.8391**	0.8489**	0.5872**	1.0000	0.6070**
CH ₂ Br ₂	0.6266**	0.6430**	0.6872**	0.6070**	1.0000

Number of replicates (n) = 63, ** indicates significance level (p) < 0.01

(c)	CHBr ₃ ^a	CH ₃ I ^a	CHCl ₃ ^a	CHBr ₂ Cl ^a	CH ₂ Br ₂ ^a
CHBr ₃ ^b	0.8390**	0.5278**	0.4296*	0.6061**	0.4269**
CH ₃ I ^b	0.8018**	0.8969**	0.5593*	0.7816**	0.6087**
CHCl ₃ ^b	0.5228**	0.4419**	0.9511*	0.4412**	0.5715**
CHBr ₂ Cl	0.6152**	0.5217**	0.3628*	0.8200**	0.4114**
CH ₂ Br ₂ ^b	0.6254**	0.6003**	0.7117*	0.5977**	0.9610**

Number of replicates (n) = 126, ** indicates significance level (p) < 0.01, ^a denotes chlorophyll a-normalized compounds; ^b denotes cell density-normalized compounds

4.3.8 Axenicity of cultures

All cultures were checked by culture on nutrient agar prior to start of experiment, and shown to be free of bacterial contamination, hence the net production of halocarbons observed relative to the subtraction of the controls are ascribed to the microalgal cultures.

4.4 Experiment 4: Effects of different irradiances on halocarbon emission

4.4.1 Growth response and pH changes

The changes of chlorophyll-a and cell density, specifically the increase of chlorophyll a and cell density as observed in Figure 4.16 and Figure 4.17 shown in percentage, were used to normalize with the emission before and after 12-hour light-exposure, which will be shown in this chapter later, in order to determine the emission rates of halocarbons from the three tropical marine microalgae.

With 40 μ mol photons m⁻² s⁻¹ serving as the control of the irradiance experiments, chlorophyll-a decreased when exposed to higher irradiance (120 μ mol photons m⁻² s⁻¹) for *Synechococcus* sp. UMACC 371 and *Parachlorella* sp. UMACC 245 as compared to *Amphora* sp. UMACC 370 (Figure 4.16) that increased. The opposite was observed when measured with cell density (Figure 4.17).

The percentage change in chlorophyll a and cell density decreased for *Amphora* sp. UMACC 370 and *Parachlorella* sp. but increased for *Synechococcus* sp. UMACC 371 when exposed to complete darkness (0 μ mol photons m⁻² s⁻¹) as compared to the control experiment (40 μ mol photons m⁻² s⁻¹).



Figure 4.16: Changes (%) in chlorophyll-a before and after 12-hour of lightexposure of the three microalgae under three different irradiance levels.



Figure 4.17: Changes (%) in cell density before and after 12-hour light-exposure of the three microalgae under three different irradiance levels.

To further assess the growth of the microalgae for all three different irradiances, other growth parameters were taken into account and recorded such as dry weight (Figure 4.18), Optical Density, OD_{620nm} (Figure 4.19), and carotenoid content (Figure 4.20).

The increase or decrease of the biomass varied amongst different algal species. For instance, no more than 15% of a change in biomass was observed for dry weight, except a large percentage increase for *Parachlorella* sp. UMACC 245 from 0 μ mol photons m⁻²

s⁻¹ to 120 μ mol photons m⁻² s⁻¹, and large decrease and increase of other biomasses ranging as low as 2% to more than 100% in chlorophyll-a, cell density and OD_{620nm} and carotenoids.



Figure 4.18: Changes (%) in dry weight before and after 12-hour light-exposure of the three microalgae under three different irradiance levels.



Figure 4.19: Changes (70) In OD_{620nm} before and after 12-nour light-exposure of the three microalgae under different irradiance levels.



Figure 4.20: Changes (%) in carotenoids before and after 12-hour light-exposure of the three microalgae under different irradiance levels.

The pH of the culture medium was measured in triplicates as pH before and after the light-exposure as shown in Figure 4.21 to allow interpretation of cell physiological changes with its surroundings with regards to halocarbon emissions. All three microalgae in general showed a decrease in pH after exposure to 12-hour of complete darkness and an increase in pH after exposure to higher irradiance (120 μ mol photons m⁻² s⁻¹) for the same amount of time. When exposed to the control level of irradiance (40 μ mol photons m⁻² s⁻¹), all three microalgae also showed an increase in pH, though the increased pH in 120 μ mol photons m⁻² s⁻¹ was higher than the control irradiance level.



Figure 4.21: Changes of pH before (dashed lines) and after (solid lines) 12-hour light-exposure of the three microalgae under three different irradiance levels. n = 3.

4.4.2 Changes of Fv/Fm as algal cell stress indicator

The maximum quantum yield, F_v/F_m , was measured as an indication of physiological state (health) of the cells (Hughes *et al.* 2006; Keng *et al.*, 2013). It is thus a good biological tool that can be used to explain the relationship between cell stress and halocarbon emission or production.

Figure 4.22 shows the F_v/F_m measured before and after 12-hour light-exposure acclimatization for all three microalgae in triplicates. A decrease in F_v/F_m after 12 hour of acclimatization under all three different irradiances, 0, 40 and 120 µmol photons m⁻² s⁻¹ was observed for all microalgae, except an increase by *Amphora* sp. UMACC 370 at 0 µmol photons m⁻² s⁻¹ and 40 µmol photons m⁻² s⁻¹. At 120 µmol photons m⁻² s⁻¹, the decrease in F_v/F_m was larger as compared to other irradiances.



Figure 4.22: Changes of maximum quantum yield, F_v/F_m , before (dashed lines) and after (solid lines) 12-hour light-exposure under three different irradiance levels for the three microalgae. n = 3.

When compared amongst taxa using mean F_v/F_m as shown in Table 4.8, there was a decrease in F_v/F_m across all three microalgae when comparing the control irradiance (40 µmol photons m⁻² s⁻¹) and the higher irradiance (120 µmol photons m⁻² s⁻¹). When cultures were exposed to complete darkness, *Synechococcus* sp. UMACC 371 and *Parachlorella* sp. UMACC 245 showed an increase in F_v/F_m , 0.32 to 0.334 and 0.64 to 0.66 respectively while *Amphora* sp. UMACC 370 showed a decrease in F_v/F_m , from 0.576 to 0.558.

Table 4.8: Mean $F_v/F_m \pm S.D.$ values of the microalgae measured under different irradiance levels. Data was statistically analyzed using Factorial ANOVA. Data presented are mean values of F_v/F_m from a total of 36 replicates (n = 36). Different letters denote standard deviation (S.D.) homogenous group (p < 0.01) according to post hoc Tukey's test.

Irradiance	Synechococcus	Parachlorella sp.	Amphora sp.
(µmol photons	sp. UMACC 371	UMACC 245	UMACC 370
$m^{-2} s^{-1}$)			
0	0.334 (±0.021) ^b	0.658(±0.020) ^{c,d}	0.558(±0.010) ^{c,d}
40	0.320 (±0.007) ^b	0.640 (±0.021) ^c	0.576(±0.037) ^{a,d}
120	0.288 (±0.029) ^b	$0.592(\pm 0.087)^{a,c,d}$	$0.522(\pm 0.094)^{a}$

4.4.3 Comparison of halocarbon emissions amongst microalgae

Data acquired for the five halocarbon emission were used to normalize to biomass, chlorophyll-a and cell density to obtain emission rates. Figures 4.23 to 4.25 (chlorophyll-a normalized) and Figure 4.26 to 4.28 (cell density normalized) show the increase and decrease (indicated by positive and negative values respectively) of five halocarbon emission rates in percentage (%) after 12-hour exposure to different irradiances. This was assessed and grouped by taxa.

4.4.3.1 Normalization to chlorophyll-a

In Figure 4.23, *Synechococcus* sp. UMACC 371 showed an increase in emission rate for CH₃I, CHBr₃ and CH₂Br₂ when exposed to both complete darkness (0 μ mol photons m⁻² s⁻¹) and higher irradiance (120 μ mol photons m⁻² s⁻¹) as compared to the control irradiance (40 μ mol photons m⁻² s⁻¹). CHBr₂Cl emission rate increased when exposed to 0 and 120 μ mol photons m⁻² s⁻¹ as compared to 40 μ mol photons m⁻² s⁻¹. The emission rate of CHBr₃ and CH₂Br₂ decreased when exposed to 40 μ mol photons m⁻² s⁻¹.

The emission rates of all five compounds, CHBr₃, CH₃I, CHCl₃, CHBr₂Cl and CH₂Br₂ for *Parachlorella* sp. UMACC 245 as shown in Figure 4.24 decreased when exposed to 40 μ mol photons m⁻² s⁻¹. All compounds except CH₃I showed an increase in emission rates when exposed to 120 μ mol photons m⁻² s⁻¹. All compounds except CH₂Br₂ showed an increase in emission rate when exposed to 0 μ mol photons m⁻² s⁻¹.

CHBr₃ emitted by *Amphora* sp. UMACC 370 in Figure 4.25 showed a decrease in emission rates when exposed to 40 and 120 μ mol photons m⁻² s⁻¹. When exposed to 120 μ mol photons m⁻² s⁻¹, the emission rate for CH₃I and CH₂Br₂ increased. The emission rate of CHBr₂Cl remained relatively unchanged when exposed to 120 μ mol photons m⁻² s⁻¹. When exposed to 0 and 120 μ mol photons m⁻² s⁻¹ as compared to 40 μ mol photons m⁻² s⁻¹.

showed a decrease in emission when exposed to 120 μ mol photons m⁻² s⁻¹. All five compounds showed an increase in emission rates when exposed to 0 μ mol photons m⁻² s⁻¹.



Figure 4.23: Percent changes of the five halocarbon emission rates normalized to chlorophyll-a by *Synechococcus* sp. UMACC 371 under three different irradiance levels, 0, 40, 120 μmol photons.m⁻² s⁻¹.



Figure 4.24: Percent changes of the five halocarbon emission rates normalized to chlorophyll-a by *Parachlorella* sp. UMACC 245 under three different irradiance levels, 0, 40, 120 μmol photons.m⁻² s⁻¹.


Figure 4.25: Percent changes of the five halocarbon emission rates normalized to chlorophyll-a by *Amphora* sp. UMACC 370 under three different irradiance levels, 0, 40, 120 μmol photons.m⁻² s⁻¹.

4.4.3.2 Normalization to cell density

For *Synechococcus* sp. UMACC 371, the increase and decrease of emission rates for CHBr₃, CHCl₃, CHBr₂Cl and CH₂Br₂ normalized to cell density as observed in Figure 4.26 were similar to the emission rates normalized to chlorophyll-a. The emission rates for all five short-lived halocarbons normalized to cell density as observed in Figure 4.27 was consistent with the emission rates normalized to chlorophyll-a by *Parachlorella* sp. UMACC 245.

In Figure 4.28, the emission rate for all five compounds by *Amphora* sp. UMACC 370 increased when exposed to 0 μ mol photons m⁻² s⁻¹. When exposed to 120 μ mol photons m⁻² s⁻¹, the emission rate of CH₃I, CHBr₂Cl and CH₂Br₂ increased. The emission rate of CH₃I increased when exposed to both 0 and 120 μ mol photons m⁻² s⁻¹ as compared to 40 μ mol photons m⁻² s⁻¹.



Figure 4.26: Percent changes of the five halocarbon emission rates normalized to cell density by *Synechococcus* sp. UMACC 371 under three different irradiance levels, 0, 40, 120 μmol photons.m⁻² s⁻¹.



Figure 4.27: Percent changes of the five halocarbon emission rates normalized to cell density by *Parachlorella* sp. UMACC 245 under three different irradiance levels, 0, 40, 120 µmol photons.m⁻² s⁻¹.



Figure 4.28: Percent changes of the five halocarbon emission rates normalized to cell density by *Amphora* sp. UMACC 370 under three different irradiance levels, 0, 40, 120 μ mol photons.m⁻² s⁻¹.

4.4.4 Comparisons between irradiance and F_v/F_m amongst microalgae

Figure 4.29 shows significant (p<0.05) decrease and increase in F_v/F_m of the three microalgae after 12-hour of irradiance exposure to 0, 40 and 120 µmol photons m⁻² s⁻¹. Cultures that were exposed to higher irradiance (120 µmol photons m⁻² s⁻¹) as compared to control irradiance (40 µmol photons m⁻² s⁻¹) showed a decrease (negative values) in the F_v/F_m across all three microalgae. *Parachlorella* sp. UMACC 245 showed a significant (p<0.05) decrease, indicating the largest decrease in F_v/F_m as compared to the other two taxa. After exposure to 12-hour of complete darkness, the F_v/F_m for *Synechococcus* sp. UMACC 371 and *Amphora* sp. UMACC 370 increased but decreased for *Parachlorella* sp. UMACC 245.

Table 4.9 shows multiple comparisons of the effect of irradiance on the changes in F_v/F_m for each microalgae. The changes in F_v/F_m were significant (*p*<0.05) for all pairwise (Bonferroni) interactions at all levels of irradiance by the three microalgae, except the comparison of F_v/F_m between irradiance 0 and 40 µmol photons m⁻² s⁻¹ by *Parachlorella* sp. UMACC 245.



Figure 4.29: Changes of maximum quantum yield, F_v/F_m , across three different irradiance levels (0, 40, 120 µmol photons m⁻² s⁻¹) by the three microalgae. n =9.

Table 4.9: Pairwise comparisons of the F_v/F_m at different irradiances (0, 40, 120)
upper upper $m^{-2} s^{-1}$ amongst the three microalgae. $n = 27$.

	Pairwise Comparisons						
	Deper	ndent Var	iable: F _v /F _m				
			Mean				
Taxa	(I)	(J)	Difference	Std.	Sig. ^b		
	Light	Light	(I - J)	Error			
Parachlorella	0	40	008	.005	.447		
sp. UMACC 245		120	.131*	.005	.000		
	40	0	.008	.005	.447		
		120	.139*	.005	.000		
	120	0	131*	.005	.000		
		40	139*	.005	.000		
Amphora sp.	0	40	.048*	.005	.000		
UMACC 370		120	.159*	.005	.000		
	40	0	048*	.005	.000		
		120	.111*	.005	.000		
	120	0	159*	.005	.000		
		40	111*	.005	.000		
Synechococcus	0	40	.038*	.005	.000		
sp. UMACC 371		120	.091*	.005	.000		
	40	0	038*	.005	.000		
		120	.054*	.005	.000		
	120	0	-0.91*	.005	.000		
		40	-0.54*	.005	.000		

*. The mean difference is significant at the 0.05 level.

4.4.5 Correlation of halocarbon emission rates

The correlation of the emission rates for all five compounds are summarized in Table 4.10 based on chlorophyll-a, cell density and both, respectively. Emission rates normalized to chlorophyll-a were highly-correlated (p<0.01) for all compounds. Emission rates normalized to cell density were non-significant. When pooled together to correlate the emission rates normalized to cell density and chlorophyll-a, CH₃I, CHCl₃ and CH₂Br₂ were highly positive correlated (p<0.01), CHBr₂Cl was correlated (p<0.05) while CHBr₃ was not significantly (p<0.05) correlated. The correlation implies that emission rates normalized to chlorophyll a were independent of those normalized to cell density.

Table 4.10: Correlation of the halocarbons. Pearson Product-Moment correlation coefficient (r) of the emission rate from the five detected compounds in term of (a) chlorophyll-a, (b) cell density, (c) chlorophyll-a and cell density produced by the three microalgae with irradiance.

(a)	CHBr ₃	CH ₃ I	CHCl ₃	CH ₂ Br ₂	CHBr ₂ Cl
CHBr ₃	1.000	0.621**	0.698**	0.846**	0.808**
CH ₃ I	0.621**	1.000	0.456*	0.650**	0.579**
CHCl ₃	0.698**	0.456*	1.000	0.615**	0.694**
CH ₂ Br ₂	0.846**	0.650**	0.615**	1.000	0.678**
CHBr ₂ Cl	0.808**	0.579**	0.694**	0.678**	1.000

** indicates significance level (p) < 0.01; * = (p) < 0.05. n = 27

(b)	CHBr ₃	CH ₃ I	CHCl ₃	CH ₂ Br ₂	CHBr ₂ Cl
CHBr ₃	1.000	-0.166 ^{NS}	0.059 ^{NS}	-0.152^{NS}	-0.063 ^{NS}
CH ₃ I	-0.166^{NS}	1.000	0.317 ^{NS}	0.488**	0.804**
CHCl ₃	0.059 ^{NS}	0.317 ^{NS}	1.000	0.225^{NS}	0.278^{NS}
CH ₂ Br ₂	-0.152^{NS}	0.488**	0.225^{NS}	1.000	0.308 ^{NS}
CHBr ₂ Cl	-0.063 ^{NS}	0.804**	0.278^{NS}	0.308 ^{NS}	1.000

NS Non-significant emission rate data were pooled from triplicates for three irradiance levels. ** indicates significance level (p) < 0.01. n = 27

(c)	CHBr ₃ ^a	CH ₃ I ^a	CHCl ₃ ^a	CH ₂ Br ₂ ^a	CHBr ₂ Cl ^a
CHBr ₃ ^b	0.379 ^{NS}	0.160 ^{NS}	$0.255^{\rm NS}$	0.470*	0.340 ^{NS}
CH ₃ I ^b	0.051 ^{NS}	0.548**	$0.339^{\rm NS}$	0.715**	0.550**
CHCl ₃ ^b	$0.028^{\rm NS}$	0.161 ^{NS}	0.621*	0.360 ^{NS}	0.312^{NS}
CH ₂ Br ₂ ^b	$0.047^{\rm NS}$	0.125^{NS}	$0.157^{\rm NS}$	0.712**	0.225^{NS}
CHBr ₂ Cl ^b	0.080 ^{NS}	0.125 ^{NS}	0.279^{NS}	0.371 ^{NS}	0.419*

NS Non-significant emission rate data were pooled from triplicates for three irradiance levels. ** indicates significance level (p) < 0.01; * = (p) < 0.05. n = 54. ^a denotes cell density-normalized compounds; ^b denotes chlorophyll a-normalized compounds

4.4.6 Pairwise comparisons between halocarbon emission rates and irradiances amongst microalgae

Pairwise comparisons showed a specific relationship of significant (p<0.05) changes in halocarbon emission rates compared amongst different irradiances for every taxa, as well as the significant (p<0.05) changes in halocarbon emission rates compared amongst different taxa for each irradiance level, chlorophyll-a and cell density-normalized emissions.

4.4.6.1 Normalization to chlorophyll-a

The pairwise comparisons of changes in halocarbon emission rates normalized to chlorophyll-a are summarized by compound, namely CHBr₃ (Figure 4.30), CH₃I (Figure 4.31), CHCl₃ (Figure 4.32), CH₂Br₂ (Figure 4.33) and CHBr₂Cl (Figure 4.34) for all three microalgae under three irradiances. Pairwise comparisons with significant (p<0.05) changes in halocarbon emission rates involving two factors, irradiances and taxa, are summarized in tables based on all five compounds, namely CHBr₃ (Table 4.11), CH₃I (Table 4.12), CHCl₃ (Table 4.13), CH₂Br₂ (Table 4.14) and CHBr₂Cl (Table 4.15). The increase and decrease of halocarbon emission rates after 12-hour exposure to different irradiances for all three microalgae are indicated with positive and negative values in figures and tables.

Figure 4.30 and Table 4.11 showed significant (p<0.05) changes in halocarbon emission rates at 0, 40 and 120 µmol photons m⁻² s⁻¹ for *Parachorella* sp. UMACC 245 but not *Synechococcus* sp. UMACC 371 and *Amphora* sp. UMACC 370. All three taxa showed significant (p<0.05) changes in CH₃I emission rates when exposed to all three different irradiances. The emission rates of CH₃I at 40 µmol photons m⁻² s⁻¹ as shown in Figure 4.31 and Table 4.12 were the lowest as compared to other irradiances. In Figure 4.32 and Table 4.13, significant (p<0.05) changes in emission rates for CHCl₃ were observed for *Parachlorella* sp. UMACC 245 when exposed to all three irradiances, but not *Synechococcus* sp. UMACC 371 and *Amphora* sp. UMACC 370. Significant (p<0.05) changes in CH₂Br₂ emission rates (Figure 4.33 and Table 4.14) were observed across all three different irradiances for *Parachlorella* sp. UMACC 245 and *Synechococcous* sp. UMACC 2371. Significant (p<0.05) changes in CHBr₂Cl emission rates (Figure 4.34 and Table 4.15) were observed only for *Parachlorella* sp. UMACC 245 across all irradiances except the changes between 0 and 120 µmol photons m⁻² s⁻¹.

Parachlorella sp. UMACC 245 showed the biggest change in emissions of all five halocarbons as compared to *Synechococcus* sp. UMACC 371 and *Amphora* sp. UMACC 370 under 120 μmol photons m⁻² s⁻¹ and chlorophyll-a normalized condition. CH₃I showed the biggest in emission changes as compared to other halocarbons namely, CHBr₃, CHCl₃, CH₂Br₂ and CHBr₂Cl, while CH₂Br₂ was shown to be the second largest.



Figure 4.30: Changes of CHBr₃ emission rates normalized to chlorophyll-a from the microalgae under three irradiance levels (0, 40, 120 µmol photons $m^{-2} s^{-1}$). n= 27

Table 4.11: Pairwise comparisons of CHB	r3 at different irradiances (0, 40, 120 umol
photons m ⁻² s ⁻¹) amongst the	three microalgae. $n = 27$.

Pairwise Comparisons							
	Dependent Variable: CHBr ₃						
			Mean				
Taxa	(I)	(J)	Difference	Std.	Sig. ^b		
	Light	Light	(I – J)	Error			
Parachlorella	0	40	1.918*	.199	.000		
sp.		120	004	.199	1.000		
UMACC 245	40	0	-1.918*	.199	.000		
		120	-1.922*	.199	.000		
	120	0	.004	.199	1.000		
		40	1.922*	.199	.000		
Amphora sp.	0	40	.045	.199	1.000		
UMACC 370		120	.418	.199	.150		
	40	0	045	.199	1.000		
		120	.373	.199	.231		
	120	0	418	.199	.150		
		40	373	.199	.231		
Synechococcus	0	40	.384	.199	.208		
sp. UMACC 371		120	.000	.199	1.000		
	40	0	384	.199	.208		
		120	384	.199	.208		
	120	0	000	.199	1.000		
		40	384	.199	.208		

Based on estimated marginal means.

*. The mean difference is significant at the 0.05 level.



Figure 4.31: Changes of CH₃I emission rates normalized to chlorophyll-a from microalgae under three irradiance levels (0, 40, 120 μ mol photons m⁻² s⁻¹). n= 27.

	Pairwise Comparisons					
	Deper	ndent Var	iable: CH ₃ I			
			Mean			
Taxa	(I)	(J)	Difference	Std.	Sig. ^b	
	Light	Light	(I - J)	Error		
Parachlorella	0	40	1 086*	067	000	
r uruchiorellu	0	120	7/8*	.007	.000	
IIMACC 245	40	0	1.086*	.007	.000	
OWNREE 245	40	120	-1.000	.007	.000	
	120	0	558	.007	.000	
	120	40	/40	.007	.000	
1mphora op	0	40	.558	.007	.000	
IMACC 370	0	120	.913*	.007	.000	
UNIACC 570	40	120	.574	.007	.000	
	40	120	915	.007	.000	
	120	120	341	.007	.000	
	120	0 40	$3/4^{\circ}$.007	.000	
Course la sessione	0	40	341*	.007	.000	
Synechococcus	0	40	.394*	.007	.000	
sp. UMACC 3/1		120	.201*	.067	.003	
	40	0	594*	.067	.000	
		120	333*	.067	.000	
	120	0	261*	.067	.003	
		40	.333*	.067	.000	

Table 4.12: Pairwise comparisons of CH_3I at different irradiances (0, 40, 120 µmol photons m⁻² s⁻¹) amongst the three microalgae. n = 27.

*. The mean difference is significant at the 0.05 level.



Figure 4.32: Changes of CHCl₃ emission rates normalized to chlorophyll-a from microalgae under three irradiance levels (0, 40, 120 µmol photons $m^{-2} s^{-1}$). n= 27.

Table 4.13: Pairwise com	parisons of CHCl ₃ a	t different irradia	(0, 40, 10)	120 ymol
photons m ⁻²	² s ⁻¹) amongst the thr	ee microalgae. n	= 27.	

Pairwise Comparisons					
	Deper	ndent Vari	able: CHCl ₃		
_			Mean	~ .	ar h
Taxa	(1)	(J)	Difference	Std.	Sig. ⁰
	Light	Light	(I – J)	Error	
Parachlorella	0	40	2.089*	.639	.013
sp.		120	.015	.639	1.000
UMACC 245	40	0	-2.089*	.639	.013
		120	-2.074*	.639	.013
	120	0	015	.639	1.000
		40	2.074*	.639	.013
Amphora sp.	0	40	019	.639	1.000
UMACC 370		120	.074	.639	1.000
	40	0	.019	.639	1.000
		120	.093	.639	1.000
	120	0	074	.639	1.000
		40	093	.639	1.000
Synechococcus	0	40	.354	.639	1.000
sp. UMACC 371		120	.020	.639	1.000
	40	0	354	.639	1.000
		120	334	.639	1.000
	120	0	020	.639	1.000
		40	.334	.639	1.000

Based on estimated marginal means.

*. The mean difference is significant at the 0.05 level.



Figure 4.33: Changes of CH₂Br₂ emission rates normalized to chlorophyll-a from microalgae under three irradiance levels (0, 40, 120 µmol photons m⁻² s⁻¹). n= 27. Table 4.14: Pairwise comparisons of CH₂Br₂ at different irradiances (0, 40, 120 µmol photons m⁻² s⁻¹) amongst the three microalgae. n = 27.

	-	• • ~			
Pairwise Comparisons					
	Depen	dent Varia	able: CH ₂ Br ₂		
			Mean		
Taxa	(I)	(J)	Difference	Std.	Sig. ^b
	Light	Light	(I – J)	Error	
Parachlorella	0	40	659*	031	000
sp.		120	123*	.031	.002
ÚMACC 245	40	0	659*	.031	.000
		120	782*	.031	.000
	120	0	.123*	.031	.002
		40	.782*	.031	.000
Amphora sp.	0	40	.012	.031	1.000
UMACC 370		120	.001	.031	1.000
	40	0	012	.031	1.000
		120	011	.031	1.000
)	120	0	001	.031	1.000
		40	.011	.031	1.000
Synechococcus	0	40	.543*	.031	.000
sp. UMACC 371		120	.062	.031	.179
	40	0	543*	.031	.000
		120	481*	.031	.000
	120	0	062	.031	.179
		40	.481*	.031	.000

*. The mean difference is significant at the 0.05 level.





Table 4.15: Pairwise co	mparisons of CH	IBr ₂ Cl at differen	t irradiances (0), 40, 120	0
ųmol photons	m ⁻² s ⁻¹) amongst	the three microal	gae. $n = 27$.		

Pairwise Comparisons						
	Depend	lent Varia	ble: CHBr ₂ Cl			
			Mean			
Taxa	(I)	(J)	Difference	Std.	Sig. ^b	
	Light	Light	(I – J)	Error		
Parachlorella	0	40	017*	002	000	
sp.		120	002	.002	1.000	
UMACC 245	40	0	017*	.002	.000	
		120	015*	.002	.000	
	120	0	002	.002	1.000	
		40	.015*	.002	.000	
Amphora sp.	0	40	.000	.002	1.000	
UMACC 370		120	.001	.002	1.000	
	40	0	.000	.002	1.000	
		120	.001	.002	1.000	
	120	0	001	.002	1.000	
		40	001	.002	1.000	
Synechococcus	0	40	001	.002	1.000	
sp. UMACC 371		120	001	.002	1.000	
	40	0	.001	.002	1.000	
		120	.000	.002	1.000	
	120	0	.001	.002	1.000	
		40	.000	.002	1.000	

*. The mean difference is significant at the 0.05 level.

4.4.6.2 Normalization to cell density

The pairwise comparisons of changes in halocarbon emission rates normalized to cell density are summarized by compound, namely CHBr₃ (Figure 4.35), CH₃I (Figure 4.36), CHCl₃ (Figure 4.37), CH₂Br₂ (Figure 4.38) and CHBr₂Cl (Figure 4.39) for all three microalgae under all three irradiances. Pairwise comparisons with significant (p<0.05) changes in halocarbon emission rates involving the two factors, irradiances and taxa, are summarized in tables based on all five compounds, namely CHBr₃ (Table 4.16), CH₃I (Table 4.17), CHCl₃ (Table 4.18), CH₂Br₂ (Table 4.19) and CHBr₂Cl (Table 4.20). The increase and decrease of halocarbon emission rates after 12-hour exposure to different irradiances for all three microalgae are indicated with positive and negative values in figures and tables.

The changes in emission rates of CHBr₃ (Figure 4.35 and Table 4.16) were not significant when compared amongst all different irradiances for the three taxa. *Parachlorella* sp. UMACC 245 and *Amphora* sp. UMACC 370 showed significant (p<0.05) changes in CH₃I emission rates (Figure 4.36 and Table 4.17). The changes of CHCl₃ emission rates were not significant (Table 4.18) across all irradiances but clear increase and decrease of halocarbon emission rate changes (Figure 4.37). Unlike the rest of other compounds, CH₂Br₂ normalized to cell density as shown in Table 4.19 and Figure 4.38 showed significant (p<0.05) changes in emission rates across all different irradiances for *Parachlorella* sp. UMACC 245, *Synechococcus* sp. UMACC 371 and *Amphora* sp. UMACC 370. *Amphora* sp. UMACC 370 was the only taxa to show a significant (p<0.05) change in CHBr₂Cl emission rates across all irradiances. Note that the mean differences obtained as shown in Table 4.16 to 4.20 were corrected to the power of 10¹¹ for statistical analysis purpose and hence were not actual values.

The halocarbon emission normalized to cell density, by *Parachlorella* sp. UMAC 245 and *Synechococcus* sp. UMACC 371, was not significantly affected by exposure to higher irradiance (120 μ mol photons m⁻² s⁻¹). *Amphora* sp. UMACC 370 showed significance (p<0.05) changes in all five halocarbon emission rate.



Figure 4.35: Changes of CHBr₃ emission rates normalized to cell density from microalgae under three irradiance levels (0, 40, 120 μ mol photons m⁻² s⁻¹). n= 27.

Pairwise Comparisons							
Dependent Variable: CHBr ₃							
	Mean						
Taxa	(I)	(J)	Difference	Std.	Sig. ^b		
	Light	Light	(I – J)	Error			
Parachlorella	0	40	400 700	1040 981	1 000		
sp.	Ū	120	-1.167	1040.981	1.000		
UMACC 245	40	0	-400.700	1040.981	1.000		
		120	-401.867	1040.981	1.000		
	120	0	1.167	1040.981	1.000		
		40	401.867	1040.981	1.000		
Amphora sp.	0	40	235.767	1040.981	1.000		
UMACC 370		120	2238.600	1040.981	.136		
	40	0	-235.767	1040.981	1.000		
		120	2002.833	1040.981	.211		
	120	0	-2238.60	1040.981	.136		
		40	-2002.83	1040.981	.211		
Synechococcus	0	40	145.200	1040.981	1.000		
sp. UMACC 371		120	.000	1040.981	1.000		
	40	0	-145.200	1040.981	1.000		
		120	-145.200	1040.981	1.000		
	120	0	000	1040.981	1.000		
		40	145.200	1040.981	1.000		

Table 4.16: Pairwise comparisons of CHBr3 at different irradiances (0, 40, 120 umo	1
photons $m^{-2} s^{-1}$) amongst the three microalgae. $n = 27$.	

*. The mean difference is significant at the 0.05 level.

b. Adjustment for multiple comparisons: Bonferroni.



Figure 4.36: Changes of CH_3I emission rates normalized to cell density from microalgae under three irradiance levels (0, 40, 120 µmol photons m⁻² s⁻¹). n= 27.

Pairwise Comparisons						
	Depe	endent Va	ariable: CH ₃ I			
			Mean			
Taxa	(I)	(J)	Difference (I	Std.	Sig. ^b	
	Light	Light	– J)	Error		
Parachlorella	0	40	268.700*	97.109	.038	
sp.		120	205.033	97.109	.147	
UMACC 245	40	0	-268.700*	97.109	.038	
		120	-63.667	97.109	1.000	
	120	0	-205.033	97.109	.147	
		40	63.667	97.109	1.000	
Amphora sp.	0	40	5065.467*	97.109	.000	
UMACC 370		120	2663.933*	97.109	.000	
	40	0	-5065.47*	97.109	.000	
		120	-2401.53*	97.109	.000	
	120	0	-2663.93*	97.109	.000	
		40	2401.533*	97.109	.000	
Synechococcus	0	40	-90.733	97.109	1.000	
sp. UMACC 371		120	-23.000	97.109	1.000	
	40	0	90.733	97.109	1.000	
		120	67.733	97.109	1.000	
	120	0	23.000	97.109	1.000	
		40	-67.733	97.109	1.000	

Table 4.17: Pairwise	comparisons of CH	I ₃ I at different	irradiances (0	, 40, 120 ymol
photons	m ⁻² s ⁻¹) amongst th	e three microa	lgae. $n = 27$.	

Based on estimated marginal means.*. The mean difference is significant at the 0.05 level.b. Adjustment for multiple comparisons: Bonferroni.



Figure 4.37: Changes of CHCl₃ emission rates normalized to cell density from microalgae under three irradiance levels (0, 40, 120 μ mol photons m⁻² s⁻¹). n= 27.

Pairwise Comparisons						
	Depe	endent Va	riable: CHCl ₃			
			Mean			
Taxa	(I)	(J)	Difference	Std.	Sig. ^b	
	Light	Light	(I – J)	Error		
Parachlorella	0	40	347.000	382.989	1.000	
sp.		120	15.233	382.989	1.000	
UMACC 245	40	0	-347.000	382.989	1.000	
		120	-331.767	382.989	1.000	
-	120	0	-15.233	382.989	1.000	
		40	331.767	382.989	1.000	
Amphora sp.	0	40	751.400	382.989	.196	
UMACC 370		120	377.333	382.989	1.000	
	40	0	-751.400	382.989	.196	
		120	-374.067	382.989	.000	
	120	0	-377.333	382.989	1.000	
		40	374.067	382.989	1.000	
Synechococcus	0	40	-109.733	382.989	1.000	
sp. UMACC 371		120	-11.467	382.989	1.000	
	40	0	109.733	382.989	1.000	
		120	98.267	382.989	1.000	
	120	0	11.467	382.989	1.000	
		40	-98.267	382.989	1.000	

Table 4.18: Pairwise comparisons of CHCl ₃ at different irradiances (0, 40, 120 umo	1
photons $m^{-2} s^{-1}$) amongst the three microalgae. $n = 27$.	

*. The mean difference is significant at the 0.05 level.

b. Adjustment for multiple comparisons: Bonferroni.



Figure 4.38: Changes of CH_2Br_2 emission rates normalized to cell density from microalgae under three irradiance levels (0, 40, 120 µmol photons m⁻² s⁻¹). n= 27.

Pairwise Comparisons						
	Depe	ndent Vai	riable: CH ₂ Br ₂			
			Mean		_	
Taxa	(I)	(J)	Difference	Std.	Sig. ^b	
	Light	Light	(I – J)	Error		
Parachlorella	0	40	137.767*	382.989	1.000	
sp.		120	-11.367	382.989	1.000	
UMACC 245	40	0	-138.77*	382.989	1.000	
		120	-150.13*	382.989	1.000	
	120	0	11.367	382.989	1.000	
		40	150.133*	382.989	1.000	
Amphora sp.	0	40	217.600*	382.989	.196	
UMACC 370		120	-2.800	382.989	1.000	
	40	0	-217.60*	382.989	.196	
		120	-220.40*	382.989	.000	
	120	0	-2.800	382.989	1.000	
		40	220.400*	382.989	1.000	
Synechococcus	0	40	177.300*	382.989	1.000	
sp. UMACC 371		120	-2.967	382.989	1.000	
	40	0	-177.30*	382.989	1.000	
		120	-180.27*	382.989	1.000	
	120	0	2.967	382.989	1.000	
		40	180.267*	382.989	1.000	

Γable 4.19: Pairwise comparisons of CH ₂ Br ₂ at different irradiances (0, 40,	120 ųmo)l
photons $m^{-2} s^{-1}$) amongst the three microalgae. $n = 27$.		

*. The mean difference is significant at the 0.05 level.

b. Adjustment for multiple comparisons: Bonferroni.



Figure 4.39: Changes of CHBr₂Cl emission rates normalized to cell density from microalgae under three irradiance levels (0, 40, 120 µmol photons $m^{-2} s^{-1}$). n= 27.

Pairwise Comparisons					
	Depend	ent Varia	able: CHBr ₂ Cl		
			Mean		
Taxa	(I)	(J)	Difference	Std.	Sig. ^b
	Light	Light	(I – J)	Error	
Parachlorella	0	40	3.700	2.375	.410
sp.		120	1.033	2.375	1.000
UMACC 245	40	0	-3.700	2.375	.410
		120	-2.667	2.375	.829
	120	0	-1.033	2.375	1.000
		40	2.667	2.375	.829
Amphora sp.	0	40	9.133*	2.375	.004
UMACC 370		120	6.733*	2.375	.033
	40	0	-9.133*	2.375	.004
		120	-2.400	2.375	.977
	120	0	-6.733*	2.375	.033
		40	2.400	2.375	.977
Synechococcus	0	40	-1.833	2.375	1.000
sp. UMACC 371		120	600	2.375	1.000
	40	0	1.833	2.375	1.000
		120	1.233	2.375	1.000
	120	0	.600	2.375	1.000
		40	-1.233	2.375	1.000

Table 4.20: Pairwise comparisons of CHBr₂Cl at different irradiances (0, 40, 120 μ mol photons m⁻² s⁻¹) amongst the three microalgae. n = 27.

*. The mean difference is significant at the 0.05 level.

b. Adjustment for multiple comparisons: Bonferroni.

4.4.7 Correlation between F_v/F_m and halocarbon emission rates at different irradiances amongst microalgae

The correlation between halocarbon emission rates and maximum quantum yield, F_v/F_m , of the three microalgae under three different irradiances is summarized in Table 4.21 and Table 4.22, normalized to chlorophyll-a and cell density respectively.

Emission rate of all five compounds normalized to chlorophyll-a and cell density for *Parachlorella* sp. UMACC 245 and *Amphora* sp. UMAC 370 were negatively correlated to the changes in F_v/F_m , except CH₃I that showed the only positive correlation. Emission rates of CH₃I, CHCl₃ and CH₂Br₂ normalized to chlorophyll-a and emission rates of CHCl₃ normalized to cell density showed positive correlation to the F_v/F_m while the rest

showed negative correlations. The correlations between $F_{\nu}/F_{\rm m}$ and halocarbon emissions

were non-significant.

Table 4.21: Pearson correlation coefficient (r) between net changes of maximum quantum yields (F_v/F_m) of the microalgae and their halocarbon emission rates normalized to chlorophyll-a.

Compounds	Synechococcus	Parachlorella	Amphora sp.
	sp. UMACC 371	sp. UMACC 245	UMACC 370
CHBr ₃	-0.089 ^{NS}	542 ^{NS}	.471 ^{NS}
CH ₃ I	0.298 ^{NS}	.157 ^{NS}	.430 ^{NS}
	NG	NG	
CHCl ₃	0.017 ^{NS}	523 ^{NS}	.273 ^{NS}
	0.001NS	(52NS	ocoNS
CH ₂ Br ₂	0.001***	65310	093***
CUDr.C1	0.254NS	202NS	557NS
CHDI2CI	-0.334	383	.557

NS Non-significant emission rate data were pooled from three replicates for all F_v/F_m values; n = 27.

Table 4.22: Pearson correlation coefficient (r) between net changes of maximum quantum yields (F_v/F_m) of the microalgae and their halocarbon emission rates normalized to cell density.

Compounds	Synechococcus	Parachlorella	Amphora sp.
	sp. UMACC 371	sp. UMACC 245	UMACC 370
CHBr ₃	-0.085 ^{NS}	542 ^{NS}	.471 ^{NS}
CH ₃ I	-0.158 ^{NS}	.233 ^{NS}	.323 ^{NS}
CHCl ₃	$0.025^{ m NS}$	474 ^{NS}	.139 ^{NS}
		NC	NC
CH ₂ Br ₂	-0.113 ^{NS}	592 ^{NS}	222 ^{NS}
CLID - Cl	0.211NS	240NS	20CNS
CHBr ₂ Cl	-0.211***	240***	.380110

NS Non-significant emission rate data were pooled from three replicates for all F_v/F_m values; n = 27.

CHAPTER 5: DISCUSSION

5.1 Emission rates amongst the three tropical microalgae

5.1.1 Effect of different growth phases

The volatile short-lived halocarbons detected in the microalgae are CHBr₃, CH₃I, CHCl₃, CHBr₂Cl and CH₂Br₂. *Amphora* sp. UMACC 371 emitted significantly (p<0.05) higher concentrations of halogenated compounds, especially CH₃I as compared to *Synechococcus* sp. UMACC 371 and *Parachlorella* sp. UMACC 245. The emission of CH₃I is significantly (p<0.05) higher compared to other detected compounds, CHBr₃, CHCl₃, CHBr₂Cl and CH₂Br₂.

In this study, the estimated range of emission rates of each halocarbon that varied amongst the three microalgae suggested that the emission rates of each halogenated compound were species- dependent due to the different algal growth physiology at exponential and stationary phases. The higher emission rate for all five halocarbons during exponential phase than in stationary phase for *Synechococcus* sp. UMACC 371, *Amphora* sp. UMACC 370 and *Parachlorella* sp. UMACC 245 when normalized to chlorophyll-a (except CHBr₃ and CH₂Br₂ for the *Parachlorella*) and cell density, suggested that the emission of these halocarbons over 12 days of culturing were growth phase-dependent. None of the five halocarbons was found to be emitted in the same amount and concentration from the same microalgal species over the culture period, suggesting that the emissions of halocarbon may be compound-specific and strain-specific despite originating from the same microalgal species.

In this growth-cycle study, halocarbon emission rates were higher at exponential phase in general for the three microalgae. Exponential phase cells are actively growing and in a healthy state. As the culture proceed to stationary phase, the cell growth slows down and eventually stops due to chemical and physical changes such as nutrients,

irradiance and increase in inhibitory compounds in the medium. Consumption of the inorganic carbon source results in pH increase in the medium, which would influence algal activity. While it is often assumed that physiological stress does occur when microalgal cells transit from exponential to stationary phase due to limiting conditions and the stress would trigger haloperoxidase mechanism to produce more halocarbons (Moore et al., 1996), the present study indicates otherwise. All five halocarbons detected by the three tropical microalgae were found to emit at higher rates at exponential phases, with exception of two brominated compounds, CHBR₃ and CHBr₂Cl by Amphora sp. UMACC 370. Manley & de la Cuesta (1997) reported consistency of higher emission rates of CH₃I at exponential for the Bacillariophytes *Navicula* sp., *Nitzschia* sp., and *Porosira glacialis* and the Chrysophyte *Phaeocystis* sp. The higher emission rates at exponential phase may be explained as follows: i) the tropical microalgal species used in the present study may be more tolerant to the stress of an aging culture, and the condition did not lead to increased production of the halocarbons. This might have to do with the low "leakage" of hydrogen peroxide from the algal cells into the medium (Palenik et al., 1987; Wong et al., 2003); ii) the exponential phase cells are actively metabolizing, allowing higher rate of methylation of haloperoxidase for halocarbon production, as compared to the cells that experience limiting conditions in stationary phase. The haloenzymes at healthy state may be less susceptible to inhibition at its active site that allow higher chance of methylation to occur. This suggests that a more detailed research has to be done on relating the change in physiological cell state with varying nutrient composition such as sulfur, nitrogen, phosphate, that may affect the haloperoxidasemechanism; iii) higher concentration of cells in stationary phase produced less superoxide per cell than those with lower density (Marshall, 2002). As oxidative radicals produced in the cells mediate the oxidation of halides present in the medium (Neidleman & Geigert, 1986), this suggest a possibility that lower algal cell density as measured by chlorophylla and cell density during the exponential phase in this study, enhances the production of halocarbons and ultimately the emission rates. It has been reported that algal cells at exponential growth can be more toxic than those in stationary or late exponential phase (Tang & Gobler, 2009). The toxicity is caused by production of peroxidase and catalase that react with multiple compounds including organic hydroperoxides and lipid peroxides in cells. The enzymes can increase the rates of dismutation and decomposition reaction of other highly reactive oxidative species (ROS) into H₂O₂. Thus, H₂O₂ surge in the cells from these reactions may be the cause to trigger halocarbon production (Tang & Gobler, 2009).

In case of the exception observed for CHBR₃, CHBr₂Cl where emission rates are significantly (p < 0.05) higher at stationary phase, these brominated compounds may be preferentially produced due to the physiological cell stress created from the limiting conditions during growth transition. Previous studies have shown an overall higher emission at stationary phase for iodomethane, CH₃I (Scarratt & Moore, 1999; Smythe-Wright et al., 2006; Brownell et al., 2010; Hughes et al., 2011) and brominated compounds, CHBr₃, CH₂Br₂ and CHBr₂Cl (Tokarczyk & Moore, 1994; Moore et al., 1996) and each of these emissions is strain-specific. Nonetheless, the discrepancies of higher emission at exponential compared with stationary phase, as compared to the present study may be largely due to: 1) non-normalized biomass emission; emission for the brominated compounds and biomass such as algal cell density were calculated separately but not normalized which makes it difficult to compare with to the emission rates in this study; emission rates were calculated in some of the previous studies but it was not possible to make comparison in terms of different growth phases, and another study compared lag and exponential phases but not stationary phase. 2) the difference in method used, such as gas-phase using head-space were used in many previous studies

while water-phase using purge-and -trap system was used, 3) it could just be that the emission rates of halogenated compounds are strain-specific.

5.2 Comparison of emission rates

5.2.1 Tropical marine phytoplankton and seaweeds

In order to assess the importance of the source of CHBr₃, CH₃I, CH₂Br₂ and CHBr₂Cl from the tropical region, a comparison is made between the emission rates found in this study and those reported from tropical macroalgae by Keng *et al.*, 2013. Table 5.1 summarizes the emission rates of both the macroalgae and microalgae.

Table 5.1: Compar	rison of emission r	ate between tr	copical macroalga	ae and marine
microalgae for CHBra	3, CH ₃ I, CH ₂ Br ₂ and	nd CHBr2Cl u	nder dry-weight	normalization

Emission rate (pmol g DW ⁻¹ hr ⁻¹)						
	Tropical seaweeds (Keng <i>et al.</i> , 2013)		Tropical marine microalgae (present study)			
Compound	Sargassum binderi	Turbinaria conoides	Padina australis	Synechococcus sp. UMACC 371	Parachlorella sp. UMACC 245	Amphora sp. UMACC 370
CHBr ₃	4.7	6.5	68.7	0.79	0.28	1.83
		x 10 ³		(Day 4)	(Day 4)	(Day 10)
CH ₃ I	11.6	20.7	34.7	0.85	0.66	2.72
				(Day 2)	(Day 2)	(Day 10)
CH ₂ Br ₂	105	620	15.1	0.02	0.05	0.24
				(Day 10)	(Day 2)	(Day 6)
CHBr ₂ Cl	88.3	175	21.1	0.02	0.01	0.20
				(Day 2)	(Day 10)	(Day 8)

For the comparison to be made possible, both laboratory-based studies were done and analyzed using similar water-phased method, which was through the GCMS equipped with purge-and-trap system conducted at 12light:12dark cycle at a range of irradiance level of 40-47 μ mol photons m⁻²s⁻¹.



Figure 5.1: Dry weight (DW) of three tropical microalgae over a growth period of 12 days.

Our results, using dry-weight (DW) as shown in Figure 5.1 and converted to the same units give emission rates of the selected four compounds. Whilst our halocarbon emission rate per unit mass range from 3 to 30000 times lower than that for seaweeds reported by Keng *et al.* (2013), the importance of marine microalgae is potentially greater due to the fact that they inhabit more than 70% of the earth's surface.However, extensive data covering larger time and spatial scales, have to be available to allow extrapolation to natural populations to properly quantify the regional (tropical) significance of the marine microalgae as source of volatile halocarbons.

5.2.2 Previous related-studies from polar and temperate regions

Brownell *et al.* (2010) reported CH₃I emission by the temperate *Synechoccoccus* sp. CCMP 2370 (clone WH 8102) over the course of 27 days. The emission peaked at approximately 22-25 pmol L⁻¹ on Day 15 during its late stationary phase, with chlorophyll-a of 0.5-1.0 μ g L⁻¹. In the present study, the emission of CH₃I by our tropical *Synechococcus* sp. UMACC 371 peaked at 0.53 pmol L⁻¹ on Day 10 during its mid-stationary phase, with chlorophyll-a content at approximately 2.0 mg L⁻¹. While there was a consistency of CH₃I emissions peak during the stationary phase for both cyanobacteria

strains, the emission by Synechococcus sp. CCMP 2370 was at an order of five times higher than that from UMACC 371. The difference may be due to: i) incubation conditions where experiments done were under lower controlled temperature of 20-21°C, higher irradiance at 60-70 µmol photons m⁻² s⁻¹ and at nutrient-repleted condition as compared to this study. It is assumed that biological processes affected by constant environmental factors such as differences in temperature, irradiance and nutrients (Brownell et al., 2010) were responsible for the lower emission of CH₃I by Synechococcus sp. UMACC 371. ii) resultant physiological condition of the two cyanobacterial strains. The difference in starting cell density of inocula as well as chlorophyll-a content obtained during the same phase when maximum CH₃I emission was achieved for both studies may contribute to the variation in emission. Hughes et al. (2011) made a similar report on CH₃I emission by the temperate Synechococcus sp. CCMP 2370 grown at 22°C under light intensity of 40 uE m⁻¹ day⁻¹ for over a total of 24 days, with exponential phase starting from Day 4 to 16. The CH₃I concentration measured throughout the experiment range from 2-4 pmol L^{-1} which are close to the medium-only control, suggesting relatively low emission of the CH₃I compound despite a long exponential phase. In other words, the emission of this iodomethane from the Synechococcus of different climatic zones is clearly strain-specific.

In the present study, *Amphora* sp. UMACC 370, a Bacillariophyte, had significantly (p<0.05) higher emission and emission rates as compared to the other two taxa from the Cyanophyta and Chlorophyta. Manley & de la Cuesta (1997) also reported higher CH₃I emission in both exponential and stationary phases from Bacillariophyta, as compared to species from Chlorophyta, Chrysophyta, Cyanophyta and Dinophyta, which further supported results from the present study of higher CH₃I emission from the Bacillariophyta than Chlorophyta and Cyanophyta. *Synechococcus*, a Cyanophyta from present and previous studies (Hughes *et al.*, 2011; Brownell *et al.*, 2010, Sæmundsdottir & Matrai,

1998; Manley & de la Cuesta, 1997) has consistently been shown as a weak emitter of CH₃I; showing either low (close to control level) or no emission and brominated compound such as CH₃Br with no emission.

5.3 Emission rates amongst the five detected compounds

From the total combined halogen mass emitted as halocarbons calculated in percentage as summarized in Table 7, the emission contribution from iodine dominates over bromine and chlorine for the taxa that emit the highest (*Amphora*) and second highest (*Synechococcus*).

Table 5.2: Total mass of emitted halides. Total halogen mass emitted as halocarbons and percentage contribution to the total from bromine, chlorine and iodine. Taxa are arranged in decreasing total mass halogens emitted order.

Taxa	Total	% Br	% Cl	% I
	halogens			
	emitted (pg)			
Amphora sp. UMACC 370	5223.6	34.39	5.93	59.7
Synechococcus sp. UMACC 371	2033.9	35.43	13.40	51.17
Parachlorella sp. UMACC 245	1573.8	32.29	47.01	21.02

Halogenating enzymes, peroxidases, catalyze the formation of carbon-hydorgen bonds using halide ions, hydrogen peroxide and an organic substrate activated for electrophilic attack. Iodoperoxidases catalyze the formation of carbon iodine bonds, whereas bromoperoxidase catalyzes iodination and bromination process, and chloroperoxidases catalyze the iodination, bromination and chlorination of organic substrates (Moore *et al.*, 1996). The higher ratio of percentange of iodine to bromine and especially chlorine suggests that *Amphora* sp. UMACC 370 and *Synechococcus* sp. UMACC 371 on the cellular level may possess more iodoperoxidase than chloroperoxidase to catalyze iodide. On the contrary, *Parachlorella* sp. UMACC 245 emitted higher chloride than the other two halides suggests that higher concentration of chloroperoxidase may be present.

Whilst iodine-containing halocarbons are not directly involved in the depletion of stratospheric ozone, they do, however through photolysis, release iodine to react rapidly with ozone to form iodine oxides (IO/OIO), which will influence the tropospheric oxidizing capacity (McFiggans *et al.*, 2000). Calvert & Lindberg (2004) reported the potential influence of iodine-containing compounds on tropospheric chemistry, where small amounts of iodinated compounds that are present in polar air mass containing representative of Br₂-BrCl- trace gas mixtures do significantly enhance ozone depletion. With significant concentration of CH₃I observed in oceanic atmospheres (Calvert & Lindberg., 2004; Yamamoto *et al.*, 2001; Blake *et al.*, 1997), it is possible that the contribution of iodine from biogenic source like *Amphora* and *Synechococcus* may be significant over the tropical region. This encourages the local measurement of IO and precursor iodine-containing compounds as well as their interaction with currently acknowledged important trace gases like O₃ and BrO in the tropics for future studies and understanding.

5.4 Effect of irradiance and photosynthetic performance on halocarbon emission by selected microalgae

The data acquired through this study showed that the emission rates of five halocarbons, CHBr₃, CH₃I, CHCl₃, CH₂Br₂ and CHBr₂Cl from the three tropical marine microalgae were influenced by irradiance through short-term exposure to higher irradiance. Both ultraviolet radiation a and b were not detected where the incubation flasks were placed. The maximum quantum yield, F_v/F_m , of the algal cultures at different irradiances were shown to be significantly different (*p*<0.05). Different irradiances were shown to be significantly different (*p*<0.05). Different irradiances were shown to influence the changes in F_v/F_m and halocarbon emission rates in general. The

correlation between F_v/F_m and the emission rates of all five compounds were weak for all three taxa. The halocarbon emission rates normalized to chlorophyll-a, cell density and both were well correlated (*p*<0.01).

Halocarbon production has long been thought to be a defense mechanism in response to bacterial infection or as a chemical defense against oxidative stress. A few suggestions may be made to explain the results from the present light stress experiment. The significant (p<0.05) increase in chlorophyll-a and cell density normalized emission rate of monohalogenated compounds including CHBr₃, CHCl₃, CH₃I and CH₂Br₂ under different irradiances, may be a result of the process of methylation of the corresponding halide ions catalyzed by halide ion methyl transferase that involves S-adenosylmethionine (SAM) (Manley, 2002). The production of a poly-halogenated compound, in case of this study, CHBr₂Cl, may involve enzymatic halogenation via haloperoxidases that were able to oxidase halogen anions. (Theiler *et al.*, 1978).

Parachlorella sp. UMACC 245 was shown to be the most light sensitive based on the significant (p<0.05) changes in emission rates of all five halocarbons under different irradiances as compared to the other two algal taxa. This may indicate that *Parachlorella* sp. UMACC 245 is more sensitive to the trigger of halocarbon production when exposed to higher irradiance and during dark cycle. This clearly indicates that the effect of irradiance on halocarbon production is species-dependent. The release of halocarbons has been a form of defense mechanism to abiotic stresses and herbivaory-predators (Bravo-Linares & Mudge, 2009; Paul & Pohnert, 2011). These halocarbons may be produced from hydrogen peroxide (H₂O₂) and the sources to produce H₂O₂ include mitochondrion transport chain (Cadenas, 1989), through the Mehler reaction or pseudocyclic photophosphorylation during photosynthesis (Pedersén *et al.*, 1996) and other cellular sources. *Parachlorella* sp. UMACC 245 in the tropics may have adapted to the ability of

producing higher concentration of H_2O_2 as a defense mechanism to cope with higher irradiance and possibly better at recovery from photoinhibition than *Synechococcus* sp. UMACC 371 and *Amphora* sp. UMACC 370.

The emission of short-lived CH₃I was shown to vary the most as compared to other halocarbons namely, CHBr₃, CHCl₃, CHBr₂Cl and CH₂Br₂ when exposed to higher irradiance, especially by Amphora sp. UMACC 370. Due to the fact that other trace gases were produced, the cells might have involved haloperoxidase enzymes to mediate the halogenation of organic compounds in the medium through the breakdown of hydrogen peroxide (Butler and Walker, 1993). Other possible pathways of the CH₃I production include iodination of inorganic iodine species such as atomic iodine (I*) by methyl groups or radicals arising from photochemical reactions (Moore & Zafiriou, 1994), and the breakdown of higher molecular weight iodine-containing organics (Fenical, 1982). Although there was a low positive correlation between F_v/F_m and halocarbon emissions under different irradiances across all three taxa (except emission normalized to cell density by Synechoccocus sp. UMACC 371), the results may indicate a link between cell stress, increased cell membrane permeability and higher emission rate of CH₃I. A loss of membrane integrity increases the release of CH₃I precursors to the surrounding medium while the cellular responses to limiting conditions promote the formation of this iodinated compound. Both oxidative stress and the associated breakdown of cellular membranes could lead to enhanced CH₃I production (Hughes et al., 2011).

The emission of other short-lived compounds including CHBr₃, CHCl₃, CH₂Br₂ and CHBr₂Cl by the three taxa after exposure to higher (excessive) irradiance may be, in general, linked to the oxidative damage in the photosynthetic apparatus that results in a decrease in photosynthetic efficiency. One of the reasons for this damage may be the damage in photosystem II (PSII) caused by the oxidation of lipids, proteins and pigments

by reactive oxygen species such as singlet oxygen ($^{1}O_{2}$), H₂O₂ and the hydroxyl radical (OH) (Hughes *et al.*, 2006).

Results from this irradiance experiment also showed higher production of halocarbons, especially CH₃I, during 12 hours of complete darkness. Whilst a mechanism behind this may not be available yet, nonetheless suggestions that could be made are 1) the cells may still be able to mediate the halide precursors in the dark and thus once exposed to the first return of photons, the halocarbons are synthesized as an indication of stress. 2) Mitochondrion transport chains, one of the sources of H₂O₂ production (Cadenas, 1989), are metabolically active during the dark reaction and the transport of electrons drive the production of H₂O₂ as a by-product (Bienert *et al.*, 2006). With the availability of H₂O₂, the biosynthesis of halogenated compounds may be made possible, 3) In the dark, chloroplast and mitochondrion in photosynthetic cells are interdependence and the possible changes of their biophysical membrane, the lipid composition of the bilayers, due to the nutrient limitation may result in higher permeability for the diffusion of H₂O₂ to occur. Higher fluxes of H₂O₂ ultimately results a chance in the production of organohalogen (Hoefnagel *et al.*, 1998; Bienert *et al.*, 2006).

In general, the weak correlation between the halocarbon emission and F_v/F_m under the influence of three different irradiances implies that the halocarbon emissions in the three algae are not strongly influenced by photosynthetic performance. This may be due to the short-term exposure even at the highest irradiance. The release of CH₃I by *Parachlorella* sp. UMACC 245, CHBr₃, CH₃I, CHCl₃ and CHBr₂Cl by *Amphora* sp. UMACC 370 as well as CH₃I, CHCl₃, CH₂Br₂ (chlorophyll-a normalized) and CHCl₃ (cell density normalized) by *Synechococcus* sp. UMACC 371, was positively correlated with the F_v/F_m values. In these cases, the increase of halogenated compounds with increased F_v/F_m values

reaching a level that could cause cell stress or membrane destruction that inhibits photosynthesis and respiration in the cells. Four out of the five compounds, CHBr₃, CHCl₃, CH₂Br₂ and CBr₂Cl emitted by *Parachlorella* sp. UMACC 245, as well as brominated compounds, CHBr3 and CH2Br2 emitted by Synechococcus sp. UMACC 371 and CH₂Br₂ emitted by Amphora sp. UMACC 370 were negatively correlated with the F_v/F_m values. Nonetheless, the fluorescence-based measures of F_v/F_m that reflects the ability of Photosystem II (PSII) reaction centers to make use of the available excitation energy, based on the correlation results, does not directly explain the physiological cell stress and the formation of halocarbons might not be photosynthetically related. Controversial studies have been reported that nutrient limitation in the batch culture medium, may or may not affect the maximum quantum yield and its measurement (Parkhill et al., 2001; Cullen et al., 1992). The possible stress of the cells resulting from the short-term exposure to the higher iradiance could also be explained from the production of H₂O₂ that originated from mitochondrial respiration, where the intracellular redox state may play a role in the maintenance and production of H₂O₂, and enzymatic catalysis such as peroxisome-associated catalase (Gross, 1993) that chemically interacts with dissolved organic matter by the microalgal cells (Lin & Manley, 2012; Wever et al., 1991).

In order to give more insight on the effect of irradiance on halocarbon emissions, studies based on a longer period of exposure to the same irradiance level, and the incorporation of additional photosynthetic parameters such as the maximum Electron Transport Rate (ETR_{max}), Non-Photochemical Quenching (NPQ) and Alpha, as well as the biochemical profiling that includes fatty acid content of the cells, would be useful to fully assess the correlation between halocarbon emission and cell stress for exact explanation and source of the stress.

The emission of each halide contributed from the five detected compounds within the 12-hour light exposure in terms of total halide mass calculated in percentage is summarized in Table 5.3. From the results, *Amphora* sp. UMACC 370 was the stronger emitter of halocarbons in terms of total halogen emission (pg) as compared to the other two taxa. *Amphora* sp. UMACC 370 was shown to contribute the highest amount of iodine as compared to bromine and chlorine and even amongst the other two taxa. This was consistent with the earlier halocarbon experiment on growth-stages.

When comparing the light stress experiment to the previous growth-cycle experiment, *Parachlorella* sp. UMACC 245 showed higher ratio of iodine to bromine and chlorine as compared to the higher ratio of chlorine to bromine and iodine, respectively. *Synechococcus* sp. UMACC 371 showed a change in the ratio of three halide composition. These clearly imply a positive influence of light stress on the emission of halocarbons, despite the short stress period that lasted only for 12 hours. Cmparing the iodine release between growth stage and irradiance experiments, the percentage of iodine contributed by *Amphora* sp. UMACC 370 and *Parachlorella* sp. UMACC 245 has shown to increase from 60% to 87% and 21% to 71% respectively. Several explanations could be that these two taxa possess cell structure and size that may be more susceptible to lysis when exposed to higher irradiance, thus releasing more CH₃I (Hughes *et al.*, 2011). The possible higher concentration of iodoperoxidase present in these two taxa may also enhance the production of iodine, thought the total iodine percentage emitted by *Amphora* sp. UMACC 370 is several times higher than *Parachlorella* sp. UMACC 245.

Таха	Total	% Br	% Cl	% I
	halogens			
	emitted (pg)			
Amphora sp. UMACC 370	500.5	3.35	9.73	86.92
Synechococcus sp. UMACC 371	471.7	14.29	46.88	38.83
Parachlorella sp. UMACC 245	98.3	14.59	14.59	70.82

Table 5.3: Total mass of emitted halides. Total halogen mass emitted as halocarbons and percentage contribution to the total from bromine, chlorine and iodine. Taxa are arranged in decreasing total mass halogens emitted order.

A previous study reported higher CH₂Br₂ concentration produced by polar *Nitzschia* sp. CCMP 580 and polar Porosira glacialis CCMP 651 cultures when exposed to higher irradiance (Moore et al., 1996). In low light (12 µmol photons m⁻² s⁻¹), CH₂Br₂ emitted by Nitzschia sp. and Porosira glacialis CCMP 651 ranged from 0- 200 pmol L⁻¹ and 0-380 pmol L⁻¹ respectively. In higher light (40 µmol photons $m^{-2} s^{-1}$), the range of CH₂Br₂ emitted increased up to 1300 pmol L⁻¹ and 1600 pmol L⁻¹ respectively over a period of 30 days (Moore et al., 1996). This was consistent with the present light stress experiments as observed from Figure 4.23 to 4.28 where there was higher CH₂Br₂ concentration, or an increase in the emission when exposed to higher irradiance for Synechococcus sp. UMACC 370 and Parachlorella sp. UMACC 245, but not Amphora sp. UMACC 370, which showed no obvious changes in CH₂Br₂ concentration. Moore et al. (1996) also reported no clear trend of "higher level of illumination produces higher halocarbon concentration", while the present study had shown the significant (p < 0.05) differences in the emission rates when exposed to higher irradiance. The discrepancies between the studies could be due to 1) the use of different microalgal strains, 2) temperature used for acclimatization, 3) the amount of time exposed to higher irradiance, 4) Cultures in the previous study were exposed to the range from 12 to 40 µmol photons m⁻² s⁻¹ whereas in the present study the cultures were exposed from 40 to 120 μ mol photons m⁻² s⁻¹.

Scarratt and Moore (1999) showed a decrease in CH_3I emission by a red microalgae, *Porphyridium purpureum* despite exposure to higher irradiance (from 20 to 800 µmol photons m⁻² s ⁻¹) for a period of 24 hours. In our present study, an increase in CH_3I emission across all three taxa was observed, contrary to Scaratt and Moor's report; but this may be due to the use of different algae, thus indicating a species-specific response.

The experimental results and comparisons made were based on the laboratory incubation of microalgae instead of an *in-situ* experiment. Hence, there are still many uncertainties that may exist in the natural habitats that might further influence the emission rates of the tropical marine microalgae.

5.5 Proposed areas for future research

It should be noted that the present studies report the emissions of short-lived halocarbons by a limited number of marine tropical microalgae under a limited range of conditions. Eight compounds (others include CH₂BrI, CHBrCl₂, CH₂I₂) were initially screened while only five compounds (CH₃I, CHBr₃, CHCl₃, CH₂Br₂, CHBr₂Cl) were detected above the detection limit by GCMS to calculate for the emission (rates). More data should be collected by studies on a wide array of marine tropical microalgae and further screened for a more complete regional data of short-lived halocarbons contributed by marine microalgae from the tropics.

Our results provide the first report of halocarbon emission by monospecific marine microalgal cultures from the tropics, both on the effect of growth-stages as well as different irradiances on halocarbon emission. This contributes to the library of existing reports on halocarbon emission by phytoplankton from polar and temperate regions. Controlled studies where the algae are subjected to other environmental stress either in the laboratory or on-site, should be done for more accurate global scale normalization. Satellite-based modeling to obtain regional phytoplankton biomass such as chlorophyll-a to normalize with extrapolated data from controlled studies will be helpful to establish a direct link of exact source to the emission of the halocarbons.

Area for future research:

- To relate the effect of different environmental factors such as the changes in temperature, salinity and pH or pCO₂ have on the emissions of the marine microalgae.
- To establish a link between enzymatic mechanisms responsible for the production of halocarbons.
- To venture into molecular work that can identify the genes responsible for halocarbon production.
- 4) To further relate the process of photosynthesis through various photosynthetic parameters other than F_v/F_m to the production and emission of halocarbons.
- 5) To establish a direct link of sources of halocarbon emission in the local open ocean waters from satellite modeling observation through chlorophyll content or even ribotype abundance patterns to the halocarbon emission measured on-site and from controlled experiments in the laboratory.
- To investigate the link between oxidative stress and halocarbon production from the tropical marine microalgae.
CHAPTER 6: CONCLUSIONS

The compounds CH_3I , $CHBr_3$, $CHCl_3$, CH_2Br_2 and $CHBr_2CI$ were shown to be emitted by tropical marine microalgae, *Synechoccocus* sp. UMACC 371, *Parachlorella* sp. UMACC 245 and *Amphora* sp. UMACC 370. *Amphora* was found to have significantly (p<0.05) higher emission rates of the five short-lived halocarbons, especially CH₃I.

The emission rates for the three tropical microalgae differ between the exponential and stationary growth phases, with higher emission rates at exponential phase. Results show that the emissions of volatile short-lived halogenated compounds by the three tropical microalgae strains are not only strain-specific but also growth phase-dependent, which implies the significant role of cell growth physiological state when determining the emission rates.

The short-term exposure to a range of irradiances under controlled laboratory condition was shown to influence the emission of the five halocarbons, CH₃I, CHBr₃, CHCl₃, CH₂Br₂ and CHBr₂Cl by the three tropical marine microlagae, *Synechoccocus* sp. UMACC 371, *Parachlorella* sp. UMACC 245 and *Amphora* sp. UMACC 370. Halocarbons emitted were shown to increase and decrease when exposed to both higher irradiance and even complete darkness as compared to the control irradiance level, though the significance (p<0.05) in changes of emission for each compound that is specific to each irradiance level differ for each taxa. This clearly indicates that the effect of different irradiance on halocarbon production are species- dependent.

Parachlorella sp. UMACC 245 was shown to be the most sensitive to a change in irradiance as compared to the other two taxa based on the significant (p<0.05) changes of emission rates amongst the five halocarbons. The emission of CH₃I, especially the most abundant by *Amphora* sp. UMACC 370, was shown to be dominant in terms of their total

halide mass (pg) and significance (p < 0.05) when compared amongst the five halogenated compounds.

The different irradiances were shown to influence significantly (p<0.05) the maximum quantum yield, F_v/F_m across all three taxa; a significant (p<0.05) decrease in F_v/F_m was observed for *Parachlorella* sp. UMACC 245 and *Amphora* sp. UMACC 371 when exposed to higher irradiance at 120 µmol photons m⁻² s⁻¹ from 40 µmol photons m⁻² s⁻¹, the controlled light level. However, when exposed to complete darkness, the changes in F_v/F_m across all three taxa, despite showing increase and decrease in emission rate, did not vary significantly (p<0.05).

The changes in halocarbon emission rates were weakly correlated and not significant to the changes in F_v/F_m when compared across all three different irradiances for all three microalgae, though both positive and negative correlation specific to each compound were observed. This suggests the possibility of other physiological stress sources such as the production of H₂O₂ from not mainly photosynthesis but mitochondrion respiration, nutrient limitation and even a change of lipid composition in membrane bilayers that stimulate and drive the production of halocarbons.

The tropical marine microalgae are widely-distributed in open-water surfaces, and are amongst the most significant sources of halogen load in the stratosphere, hence there is a need to measure and monitor such emissions not only in controlled laboratory scale but also extend to look further the actual emissions under the natural environment in the tropics in the future. Studies on biological responses to climate change should transcend transient acclimation physiology to investigate long-term adaptation and evolution, besides the essential need of a comprehensive library of regional and global emission data linking compounds to algal species for the assessment and prediction of halocarbon emissions over the world-wide oceans. The research questions can be answered as follows:

- (i) The main short-lived halogenated compounds emitted by the three tropical marine microalgae, *Synechococcus* sp. UMACC 371, *Parachlorella* sp. UMACC 245 and *Amphora* sp. UMACC 370 were CHBr₃, CH₃I, CHCl₃, CHBr₂Cl and CH₂Br₂.
- (ii) The emission of short-lived halocarbons by the three tropical marine microalgae were shown to be species- and growth phase- dependent, highlighting the importance of taking cell growth physiological state when determining emission rates into consideration.
- (iii) The range of irradiances was shown to significantly (p>0.05) influence the changes in F_v/F_m and halocarbon emission rates. However, the correlation between F_v/F_m and halocarbon emission rates were weakly correlated (p>0.05), indicating that the emission of the halocarbons might not be due to algal cell stress from photosynthetic performance but may be due to other physiological mechanisms in cells.

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LIST OF PUBLICATIONS AND PAPERS PRESENTED

Yong-Kian Lim, William Sturges, Noorsaadah Abdul Rahman, Siew-Moi Phang (2016). Emission of short-lived halocarbons by selected tropical marine microalgae. Conference: Asia Oceanic Geosciences Society, Beijing, China. Date presented: 3rd August, 2016. Presentation ID: BG08-D3-AM1-311B(L3N)-006

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